

1 Assessment of the performance of river replenishment of
2 sediment applied to a non-erodible channel by means of
3 numerical modeling

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7 **Abstract**

8 The artificial replenishment of sediment is a method used to solve the
9 lack of sediment continuum in the downstream reach of dams. The un-
10 derstanding of the impact of this technique on flow and bed morphology
11 in rivers is a challenging. Furthermore it is an up-to-date topic since the
12 restoration of sediment dynamics highly impacts on biodiversity and quality
13 of the water courses. The potential of numerical schemes as a forecasting tool
14 for the decision-making processes of these replenishment procedures is still
15 poorly understood. In this work, 2D shallow water equations in combina-
16 tion with the Exner equation are solved by means of an augmented Riemann
17 solver. Computational outcomes are compared with experimental data ob-
18 tained from several replenishment configurations studied in the laboratory.
19 The classical friction approach considered for mimicking the bed channel
20 roughness has been modified for taking into account the effect of the replen-
21 ishment, which plays an important role on the temporal evolution of the
22 channel bed roughness. Results point out the promising prediction capacity
23 of the numerical solver, which can be used to evaluate currently applied or
24 to forecast the behavior of planned replenishments.

25 *Keywords:* 2D Shallow water, sediment replenishment, bed-load transport,
26 dam impact, bed-shear stress, sediment management

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27 1. Introduction

28 In the downstream reaches of dams the sediment continuum is altered
29 by the dam presence. Without sediment release structures, a dam traps the
30 sediments in the upstream reservoir, especially during floods, creating an
31 obstacle to the natural sediment transport [1, 2, 3]. Due to the absence of
32 sediment transport, many negative changes may occur along rivers on both
33 the morphological and ecological aspects. The main morphological effects
34 are related to tendency to riverbed incision, generation of an armored layer
35 and coarsening of the bed, together with bank instability and changes in
36 channel width. These changes negatively affect also the ecosystem along the
37 river, inducing a loss in the aquatic and riparian habitats with consequences
38 on the water quality [4, 5, 6, 7]. The above-mentioned outcomes limit the
39 possibilities for fish spawning and survival.

40 In the last decades, the replenishment of sediment (also called gravel aug-
41 mentation) technique has been applied in order to supply sediments lacking
42 in the downstream reaches. The artificial addition of sediments into the
43 rivers has been used since the 80' with the main purpose to recreate a nat-
44 ural sediment transport [8]. More recently, it was stated in [9, 10] that the
45 added material for replenishment purposes should be smaller than the ex-
46 isting bed material in order to favor the fine sediment availability for the
47 spawning habitats and to enhance bed elevation variations. Conversely, field
48 experiments, carried out in the United States and in several Japanese and
49 European rivers, as well as laboratory investigations have permitted to im-
50 prove the knowledge on the complex geomorphological processes occurring
51 in rivers subjected to the replenishment of sediment [11, 12]. Nowadays,
52 the replenishment of sediment is more often applied in rivers for geomor-
53 phic purposes, using a specific grain size, to maintain or even increase the
54 morphology variety including bed forms in a channel [13, 14, 12]. Extension
55 and duration of the morphological bed changes obtained by gravel augmen-
56 tation are very important in determining the effectiveness of the method [15].
57 During the last years, several laboratory experiments were performed per-
58 mitting to assess the general erosion process of deposits [7], the evolution of
59 sediment waves [16, 17, 15] and the influence of added sediment supply on
60 river bed morphology [18, 10, 19, 20]. In particular, the erosion process of
61 sediment deposits is presented by the laboratory results from [7]. Here, the
62 erosion mechanism is described as a combination of lateral erosion on the
63 sediment deposit toe by hydraulics forces, and the subsequent mass failure

64 by gravitational mechanism due to the over steepening of the banks. Field
65 observations downstream the Murou dam confirm these laboratory findings
66 [7]. After collapsing, the eroded mass of the replenishment is transported
67 further downstream. Translation and dispersion are the two main transport
68 types occurring for sediment augmentation [15, 17]. In an ideal performed
69 replenishment, as stated in [15], sediments would firstly be transported by
70 translation to the downstream target reach and, then they should shift pro-
71 gressively to a dispersion transport mechanism.

72 Beside the field and the laboratory experiments, numerical morphody-
73 namic models have been developed as a support of making-decisions about
74 the replenishment technique. [21] remarked that their numerical tool was not
75 reliable in predicting the magnitudes in bed elevation changes and failed to
76 reproduce the riffle-pool-pattern observed in field experiments with sediment
77 replenishment. In [22] the incapacity in anticipating the bed load trans-
78 port in gravel bed rivers by any of sediment transport formulas was stated.
79 Hence, the gap of knowledge about the existing link between sediment supply
80 and channel bed morphology has still to be fulfilled [16, 19]. More detailed
81 studies are required on the influencing factors affecting the success of gravel
82 augmentation projects, since the conceptional design of many of the restora-
83 tion projects are still designed based only on past experiences [5, 7, 15].
84 Above all, the role played by the discharge, the necessary amount of sedi-
85 ments and the mechanism of sediment propagation through the channel are
86 core to better understand [7, 19]. Since in-situ monitoring campaigns are
87 difficult to carry out and considering that the laboratory experiments cannot
88 fulfill all the non-yet investigated domain, the implementation of numerical
89 tools arises as a helpful approach for better understanding the morphological
90 consequences due to a sediment replenishment.

91 In this work the 2D depth-averaged mathematical model was considered
92 since this set of equations is widely accepted in the study of river configura-
93 tions and the number of equations and closure relations is limited. The mor-
94 phodynamic evolution of the replenishment material is computed by means
95 of the Exner equation. The spatial and temporal evolution of the hydrody-
96 namic and morphodynamic models must be solved through a synchronous
97 treatment, [23], since there is a strong interaction between the flow and the
98 replenishment material. There are several numerical techniques present in lit-
99 erature for solving this set of equations: Finite Volumes (FV) [24, 25, 26, 27],
100 Finite Elements (FE) [28] and Finite Differences (FD) [29]. Based on a ex-
101 plicit Finite Volume strategy described in [30] an efficient and self-stable nu-

merical scheme was validated successfully against experimental data. Hence, this numerical tool will be used for studying the hydro-morphodynamic behavior of the replenishment process when dealing with different configurations.

The remainder of the paper is structured as follows: first the methodology to obtain the experimental data is described. Then the next section is devoted to briefly outline the mathematical model and numerical scheme considered. The following section presents the results obtained with this study and their prior discussion. Finally, the conclusions of the work are summarized.

2. Experimental Methodology

Series of laboratory experiments to test different replenishment of sediment configurations were performed at the Laboratory of Hydraulic Constructions (LCH) at École Polytechnique Fédérale de Lausanne (EPFL), in Switzerland. The tests are performed in a 15 *m* long channel with a bed width of 0.4 *m*. The bank slope is 2:3 (height:length) and the longitudinal slope 0.015, Figure 1. The length of the replenishment volumes, *L*, is equal to 0.64 *m*. The discharge is controlled by a pump system. The water flows from an upper stilling basin throughout the channel reaching the outlet basin downstream. The discharge is indirectly determined by the imposed replenishment submergence condition. The replenishment height is kept constant and, in the herein presented experiments, the discharge corresponds to a completely submergence condition of the volume (volume height=water depth) equal to 19 *l/s* on model. The laboratory model has an approximate 1:10 scale factor considering the geometric characteristics of a typical alpine river. In Figure 2, the grain size distributions for the channel bed and for the replenishment material are displayed. The bed grain size distribution, used for creating the fixed bed and the banks, is representative of a typical alpine river. The replenishment volumes are composed of a finer grain distribution which dimension varying from 3 to 8 *mm*. The ecological needs for spawning grounds, scaled to laboratory scale, were considered in this grain choice [9, 10]. Each test lasts three hours. In a preliminary analysis on the evolution of the morphological bed forms demonstrated that, with this flow condition, a certain morphological equilibrium is reached in two hours, [31].

The influence of different amounts of replenishment together with four different geometrical configurations were investigated. Specifically: single volume (configuration 1), double volumes aligned (configuration 2), parallel

138 (configuration 3) and alternated (configuration 4) configurations are pro-
 139 posed, Figure 3. While in configuration 3 the volumes are symmetrically
 140 placed on the banks, in configuration 4 the two blocks on one bank are
 141 switched downstream by a distance equal to half the replenishment length.
 142 In each test the same block volume dimensions are maintained, thus the con-
 143 figurations 2, 3 and 4, with multiple volumes, have respectively two and four
 144 times the amount of gravel of configuration 1. Data analysis is performed
 145 by means of photos taken at different time steps during the tests. A camera
 146 is installed on a carriage moving along the downstream and the traversal
 147 directions, and the photos are taken each 0.5 m along the channel and then
 148 merged together forming a panoramic view. In order to facilitate the anal-
 149 ysis and the detection of the transport motion, the grains composing the
 150 replenishment were painted in red.

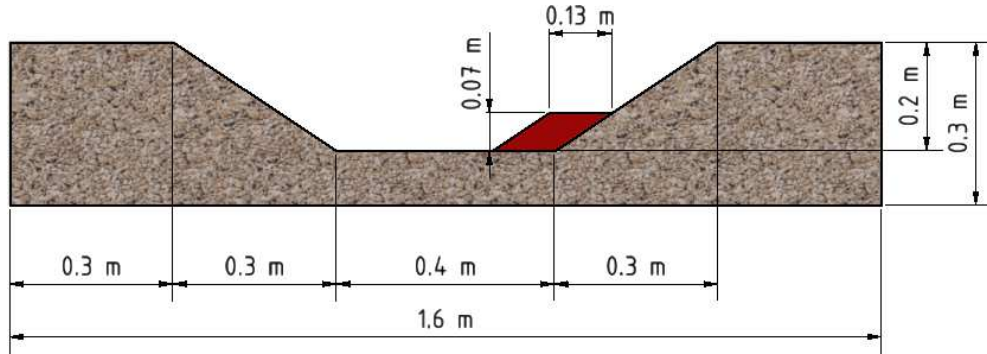


Figure 1: Cross section view of the channel with one replenishment reproduced at the bank

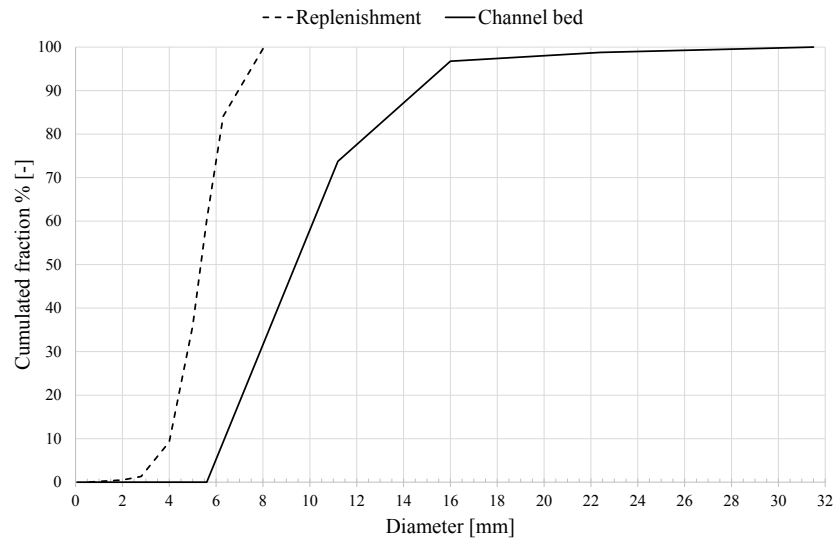


Figure 2: Grain size distribution for channel bed and replenishment material

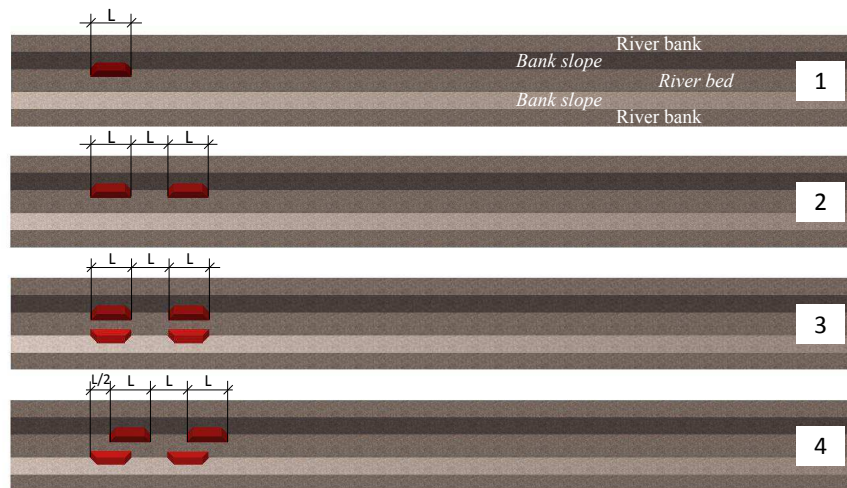


Figure 3: Volumes configurations 1, 2, 3 and 4 for the replenishment

3. Simulation Methodology

3.1. Flow simulation

The mathematical model is built by 2D shallow water equations, SWE, and the Exner equation. The SWE are derived from the Navier-Stokes equations by integrating the continuity and momentum equations over depth, [32]. The resulting 2D system of equations is written in the conservative form as follows

$$\frac{\partial \mathbf{U}}{\partial t} + \frac{\partial \mathbf{F}(\mathbf{U})}{\partial x} + \frac{\partial \mathbf{G}(\mathbf{U})}{\partial y} = \mathbf{T}_\tau + \mathbf{T}_b \quad (1)$$

where

$$\mathbf{U} = (h, hu, hv)^T \quad (2)$$

are the conserved variables with h representing water depth in the vertical coordinate and (u, v) the depth averaged components of the velocity field. The flux terms are expressed as

$$\begin{aligned} \mathbf{F} &= \left(hu, hu^2 + \frac{1}{2}gh^2, huv \right)^T \\ \mathbf{G} &= \left(hv, huv, hv^2 + \frac{1}{2}gh^2 \right)^T \end{aligned} \quad (3)$$

being g the gravity acceleration.

The source terms \mathbf{T}_τ and \mathbf{T}_b collect, respectively, the information about the friction exerted over the bed, evaluated through the 2D Manning's law, and the gravity forces written by means of the bed slopes,

$$\mathbf{T}_\tau = \left(0, -gh \frac{n^2 u \sqrt{u^2 + v^2}}{h^{4/3}}, -gh \frac{n^2 v \sqrt{u^2 + v^2}}{h^{4/3}} \right)^T \quad \mathbf{T}_b = \left(0, -gh \frac{\partial z}{\partial x}, -gh \frac{\partial z}{\partial y} \right)^T \quad (4)$$

with n the Manning's parameter and z the bed elevation.

On the other hand, the bed evolution is modeled through the Exner equation, which is a continuity equation where the bed level is modified by

the solid fluxes which cross the control volume. The 2D Exner equation for bedload is written as follows,

$$\frac{\partial z}{\partial t} + \xi \frac{\partial q_{s,x}}{\partial x} + \xi \frac{\partial q_{s,y}}{\partial y} = 0 \quad (5)$$

where $\xi = \frac{1}{1-p}$, being p the material porosity and $q_{s,x}$, $q_{s,y}$ the solid fluxes which are computed as a function of excess bed shear stress with respect to the critical value. This bedload transport is often written through the following dimensionless parameter,

$$\Phi = \frac{|\mathbf{q}_s|}{\sqrt{g(s-1)d_m^3}} \quad (6)$$

where $s = \rho_p/\rho_w$ is the ratio between solid material (ρ_p) and water (ρ_w) densities, and d_m is the median diameter of the transported sediments. According to the numerical assessment performed in [33], the Smart formula is chosen for computing the bedload discharge, which in dimensionless form is written as follows

$$\Phi = 4 \left(d_{90}/d_{30} \right)^{0.2} Fr S^{0.1} \theta^{1/2} (\theta - \theta_c^S) \quad (7)$$

where the term S is the bed slope which is evaluated as in [33] for distinguishing between in favor and adverse sloping beds, d_{90} and d_{30} are the characteristics diameters of the sample, Fr is the Froude number, θ is the dimensionless shear stress and θ_c^S is the critical shear stress according to [34].

The computational method considered for solving the equations is fully explained in [30] where the computational domain is split in fixed computational cells. Then, the Gauss theorem is applied to each volume cell for computing the solid and liquid fluxes which cross the edge of each cell. These fluxes are derived from an approximate Riemann Problem. Then, the conserved variables are updated in time and space by means of an upwind technique considering a first order approach. The stability criteria is controlled thanks to an augmented CFL condition proposed in [30].

3.2. Friction law

One of the key parameters which governs the dynamics of the sediment replenishment is the roughness of the water-worked bed. This roughness is influenced by bed mobility which has an impact on the mean longitudinal velocity and on the bed shear stress, [35, 36]. The rheological law which is

considered in this work for modeling the friction is the Manning's law, [37]. This law is based on a power-law velocity model where the friction over the bed is written as the product of a friction parameter and the square of the depth-averaged velocity, equation (4). This Manning friction parameter may be related with the grain size by means of the Strickler formula for armored channels, [38]

$$n = \frac{1}{26} d_{90}^{1/6} \quad (8)$$

Furthermore, the addition of replenishment material to the channel causes secondary effects. Due to the finer nature of the sediments of the replenishment, the initially coarser non-erodible channel suffers a change in the surface roughness, i.e. the sediments coming from the replenishment build a new layer over which the water column exerts a new friction value. In order to take into account this important characteristic, the Manning parameter is continuously updated during the simulation as follows

$$n = \begin{cases} \frac{1}{26} (d_{90}^C)^{1/6} & \text{if } DD \leq d_{90}^C + d_{90}^R \\ \frac{1}{26} (d_{90}^R)^{1/6} & \text{if } DD > d_{90}^C + d_{90}^R \end{cases} \quad (9)$$

where d_{90}^C and the d_{90}^R are the d_{90} for the non-erodible channel and the replenishment material respectively, and DD is the deposition depth of the added material for replenishment purposes.

This effect of bed fining has been confirmed in gravel-bed rivers [39, 40] and in sand-bed rivers [41, 42, 43]. For these situations, this phenomenon is a consequence of dune sorting, hydraulic structures, overbank deposition and river bifurcation. In this work, the bed fining is a consequence of the finer material of the replenishment.

4. Results & Discussion

4.1. Description of the experimental results

The 2D temporal evolution of the covered surface occupied by the replenishment material during the experiment is displayed in Figures 4, 5, 6, 7 (top views). Visual observations during the experiment show that the erosion of the blocks of replenishment occurs as a combination of fluvial erosion at the deposits toe and mass failure of the above placed sediment due to over-steepening. The collapsed mass is then further eroded and transported

226 downstream by the flow. Furthermore, in configurations 2, 3 and 4 the
 227 most upstream volumes of added material are less eroded than the follow-
 228 ing downstream replenishments. For configurations with parallel volumes,
 229 3 and 4, the deposition of the eroded replenishments occurs further down-
 230 stream. The pattern of depositions in the downstream channel varies among
 231 configurations.

232 Considering each configuration, in configuration 1 most of the material is
 233 preserved in the original location. The small amount of eroded sediments is
 234 settled in the vicinity of the channel wall. Configuration 2 (Figure 5) with
 235 two blocks in the same side leads to similar results. The material eroded from
 236 the replenishment located upstream is not able to affect the second block.
 237 Configuration 3 (Figure 6) is a step further in morphology complexity. Being
 238 the replenishment parallel aligned on the banks, in the first phases of the
 239 experience, they create a fix obstacle for the flow. In this case, the flow is
 240 forced to propagate between the blocks, i.e., in a narrower section compared
 241 to the upstream section. Thus, the flow velocity increases and the erosion
 242 rates increase as well. The above-mentioned behavior lasts until the upstream
 243 replenishment is eroded; then the flow velocity decreases together with the
 244 erosion rate, which reaches an equilibrium state. Compared with the results
 245 from configurations 1 and 2, in this case the sediments are more dispersed
 246 and more clustered depositions appear along the channel. Finally, the same
 247 behavior is observed for configuration 4 (Figure 7) which, in comparison with
 248 configuration 3 (Figure 6), shows the relevant influence of the geometrical
 249 block placement. The non-complete alignment of the replenishment upstream
 250 slows down the erosion rate and enhances the creation of a bed form pattern
 251 downstream.

252 *4.2. Comparison of the temporal evolution of the deposition evolution*

253 Numerical results are herein compared to the experimental data obtained
 254 in the laboratory for the four replenishment configurations explained in Sec-
 255 tion 2.

256 The computational domain has been discretized using a non uniform tri-
 257 angular mesh of 32000 cells, locally refined along the center area of the chan-
 258 nel. The smallest triangle area is approximately 0.0006 m^2 . At the inflow
 259 cross-section, a constant discharge equal to 19 l/s is imposed and at the
 260 outflow a free boundary condition is considered.

261 Figures 4, 5, 6, 7 (bottom views) show the temporal evolution of the
 262 computed and experimental results in terms of deposition downstream the re-

263 replenishment. The numerical results comprise 3D contours views whereas the
 264 experimental data contains 2D information obtained from the photo survey
 265 of the tests made along the time. For configuration 1, Figure 4, the material
 266 spreading covers up to 2 m downstream the block. The general tendency is
 267 well tracked by the numerical scheme. For the second configuration, Figure
 268 5, the maximum run-out (i.e. spreading) of the material is also limited and
 269 it advances 2.5 m considering the second volume of sediment. Nevertheless,
 270 the computational results are able to reproduce the phenomenon. Moreover,
 271 as visible on the experimental results, the erosion rate for both configuration
 272 1 and 2 is not complete since the completely submerging condition in those
 273 configurations is not enough for mobilizing the replenishment. The numeri-
 274 cal model predicts well the behavior, achieving a maximum deposit height,
 275 at the end of the simulation of around 0.04 m .

276 For configuration 3 (Figure 6) the computational results predict the strong
 277 mobilization of the material, which is washed away from its original location.
 278 The maximum mobilization of the material observed in the experiments, 12.5
 279 m , is almost reached with the numerical scheme. For configuration 4 (Figure
 280 7), the spreading of the material is again relevant, overtaking the channel
 281 length of 12 m . Newly, the constraint effect of the flow produced by the
 282 volumes leads to stronger erosion processes. The numerical scheme predicts
 283 successfully the maximum run-out. It is also able to reproduce the movement
 284 of the material of one of the downstream blocks which is almost completely
 285 removed and settled downstream due to the effect of the water. Conversely,
 286 although the maximum spreading and general trend is computationally well
 287 achieved, the geomorphological patterns that exhibit configurations 3 and
 288 4 are not accurately tracked. Indeed, the numerical scheme is not able to
 289 reproduce the free replenishment material areas followed by dense spatially-
 290 occupied areas.

291 Additionally, the effect of the roughness evolution of the channel due to
 292 the replenishment is herein further analyzed. For this purpose, Figure 8 dis-
 293 plays the stage at time $t=30 \text{ min}$ for the configuration 3 with a constant
 294 Manning parameter (i.e. considering only the roughness arising from the
 295 bottom material) and with a Manning parameter related with the deposition
 296 depth (i.e. assuming a shift between the roughness of the bottom and the
 297 roughness of the replenishment material). As it is observed, the overall sur-
 298 face level of the replenishment is completely different between them and also
 299 with the experimental data provided in Figure 6. An important mismatch of
 300 a physically based behavior is observed when the Manning parameter is kept

constant for all the computational domain. Therefore, the dynamic adaptation of the Manning parameter plays a key role and it has to be retained in the model. Nevertheless, it is worth noting that the larger amount of replenishment material is spread over the channel, the more relevant becomes this approach for computing the friction.

4.3. Parameter analysis

An important parameter when dealing with replenishment techniques is the occupational ratio (OR). This parameter is based on a 2D analysis and it shows the fraction of area that is covered by the replenishment material:

$$OR = \frac{A_R}{A_T} \quad (10)$$

where A_R is the area which is occupied by the replenishment material and A_T is the total area under analysis. This parameter has been computed for traversal slices of the channel with 0.1 m length. The OR aims at identifying the level of material aggregation on the channel along the longitudinal direction. For an adequate sediment replenishment process a dispersion behavior of the added material is required to achieve the most common goals to improve rivers ecology, such as fining bed grain size and enhanced bed elevation variations in order to improve rivers ecology. Hence, the OR quantifies the dispersion of the material of the replenishment along the channel.

Figure 9 shows the temporal evolution of the computational and experimental occupational ratio for each configuration tested in this work. The initial OR of the deposits of sediment are quickly increased as a consequence of the sediment transport. Sediments are washed out by the flow and the particles are spread along the channel until reaching an equilibrium stage. The time analysis of the OR shows that the position of the peaks is approximately constant in the channel. The general tendency is well reproduced by the numerical scheme, where the maximum run-out of the material is achieved for the four configurations. The values of the OR are similar over the replenishment blocks. The main differences appear in the third and fourth configurations downstream the blocks location: sediment patterns are observed in terms of a sequence of peaks of occupational ratio. In particular for configuration 4, a clear bed form pattern (transversal bars) is observed on the channel bed at the end of the experiment. The same behavior is not noticeable in the results performed by the numerical simulation. Generally, the numerical model tends to underestimated the ratio of OR. Nevertheless, the

335 difference-ratio is constant between the analyzed configurations, confirming
336 the reliability of the numerical method.

337 Travel distance (TD) of the center of mass is another important parame-
338 ter which allows to verify the geomorphological response of the channel bed
339 to gravel augmentation under given geometric and hydrodynamic conditions.
340 The higher the TD value is, the bigger the impact of the replenishment is
341 in the downstream direction. Furthermore, the analysis of the TD over time
342 allows to estimate the propagation velocity of the sediment transport. The
343 slope in the temporal evolution of transport distance can be seen as a trans-
344 port velocity, where steep slopes indicate a fast movement. Furthermore, the
345 position of the center of mass can be estimated for the entire deposition region
346 observed or partially for several percentiles. Hence, TD_x is the position where
347 the $x\%$ of the covered surface is located downstream. Figure 10 displays the
348 travel distance for several percentiles of the covered surface for the compu-
349 tational results and the experimental data. Transport velocity for TD_{75} and
350 TD_{95} are higher than of TD_{25} and TD_{25} . Each percentile travels quicker dur-
351 ing the first minutes; afterward, the transport velocity tends to slow down
352 until a equilibrium stage is reached. Moreover, for configurations 3 and 4, the
353 percentiles of the TD reach the downstream part of the channel faster due
354 to the effect of the bed fining. Regarding the numerical-experimental com-
355 parison, there is an accurate agreement between both results, although the
356 experimental travel distance is always above the computational prediction.
357 This is justified by the fact that the maximum run-out with the laboratory
358 data is slightly larger than the one obtained with the numerical tool.

359 The sediment patterns in terms of occupability ratio found in configura-
360 tion 3 and 4, as it has been mentioned before, are not accurately reproduced
361 by the numerical scheme, Figure 9. This can be justified by the fact that
362 some areas of those patterns, visible in the photo survey of the experiment,
363 are composed by a few layers of material (1-2 grains), where the friction
364 term acting on the sediment is not only governed by a basal force, but also
365 by the local collision forces among particles. With the present model, based
366 on a continuum approach, these features are not included. Further, the bed
367 roughness of the channel exhibit a chaotic nature, in the sense that although
368 a characteristic d_{90} can be estimated from the numerical simulations, the spa-
369 tial distribution of such sample is random. This fact can alter significantly
370 the patterns since they are built with 1-2 grains. However, the areas with
371 high occupation of particles are better modeled by the numerical scheme and
372 this is justified by the good fit observed in the occupational ratio and in the

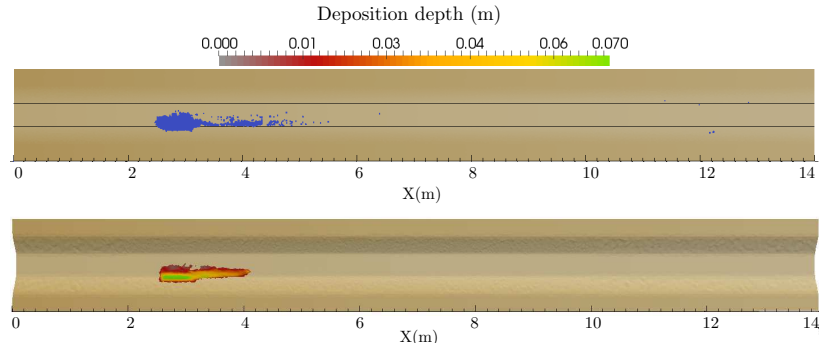
373 travel distance, Figures 9 and 10.

374 4.4. *Bed shear stress*

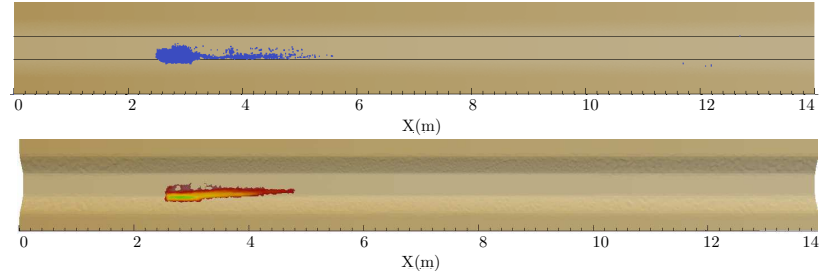
375 Regardless the configuration, the global roughness of the downstream
376 bed is reduced by the finer replenishment. Thus an increase of the bed shear
377 stress and of the transport capacity are observed. This result is in agreement
378 with the reaction of steep mountain gravel streams to a first sediment pulse
379 input, observed by [44] and with recent morphodynamic modeling, where a
380 fining of the bed surface due to increased sediment supply is predicted [45].
381 Furthermore, the decrease in roughness is in good correspondence with OR,
382 as for highest OR also the largest bed fining is observed, Figure 9.

383 For the reasons mentioned above, the estimation of the bed shear stress
384 is challenging in both field and laboratory conditions, especially with mobile
385 sediments. At the same time, knowing the value of the bed shear stress is
386 crucial for evaluating the performance of occupability ratio and travel dis-
387 tance. In this context, the numerical tool becomes helpful in forecasting the
388 results for different replenishment proposals since by means of equation (9)
389 the friction coefficient is continuously updated for each cell of the domain.
390 Figure 11 displays the temporal evolution of the bed shear stress for config-
391 uration 4. Zones having alternating higher and lower bed shear stress are
392 observed. Through numerical modeling and taking into account the replen-
393 ishment purpose (rehabilitation of the sediment continuum), it is possible to
394 determine the erosion, transport and aggradation of the sediment which is
395 fed in the channel. Hence, it will be known: (i) where to place the replen-
396 ishment, (ii) when to do it, (iii) which is the more suitable configuration of
397 blocks and (iv) if it is required to repeat the replenishment for an effective
398 final result.

Time $t = 10 \text{ min}$



Time $t = 30 \text{ min}$



Time $t = 120 \text{ min}$

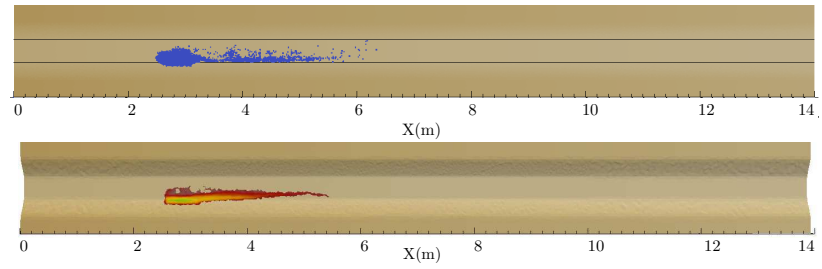
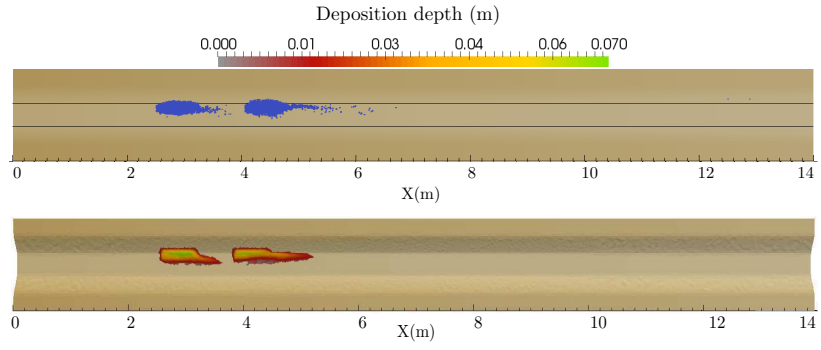
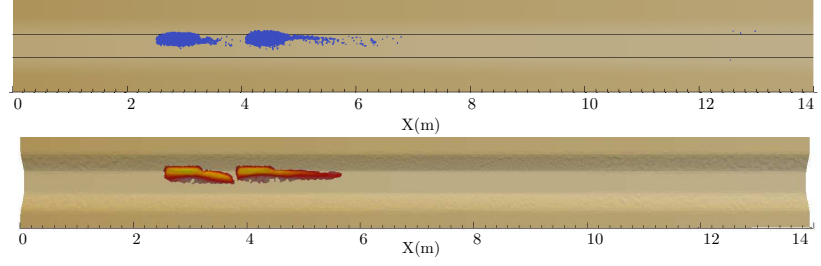


Figure 4: 2D experimental data of the covered surface (top views) and 3D computed contour views simulated numerically (bottom views) for the deposition depth along the non-erodible channel at times $t= 10, 30$ and 120 min for configuration 1

Time $t = 10 \text{ min}$



Time $t = 30 \text{ min}$



Time $t = 120 \text{ min}$

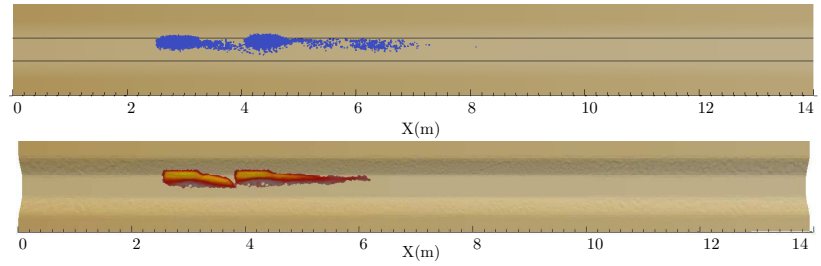
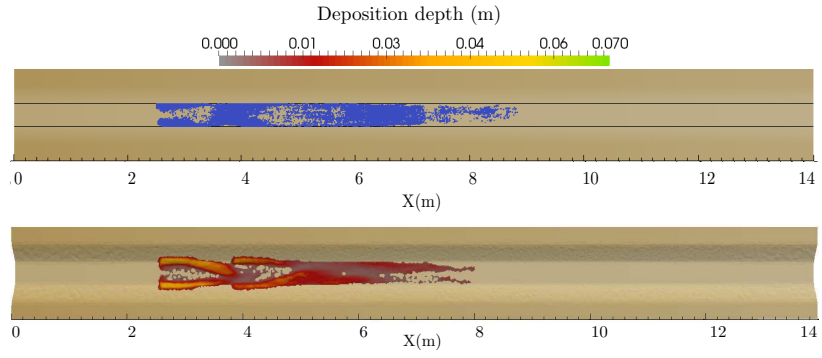
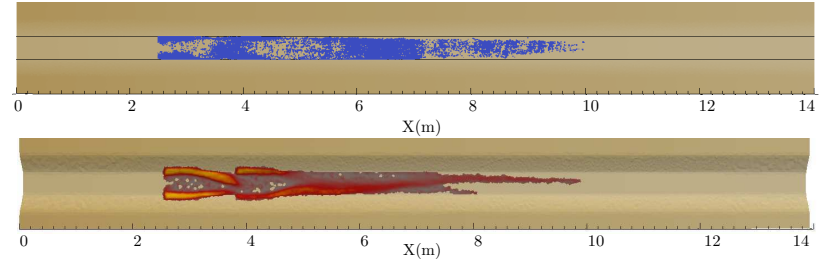


Figure 5: 2D experimental data of the covered surface (top views) and 3D computed contour views simulated numerically (bottom views) for the deposition depth along the non-erodible channel at times $t = 10, 30$ and 120 min for configuration 2

Time $t = 10 \text{ min}$



Time $t = 30 \text{ min}$



Time $t = 120 \text{ min}$

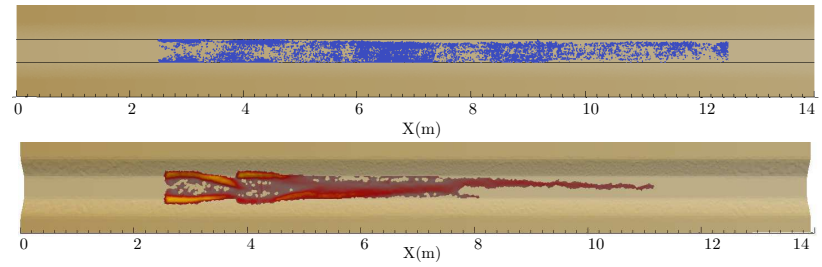
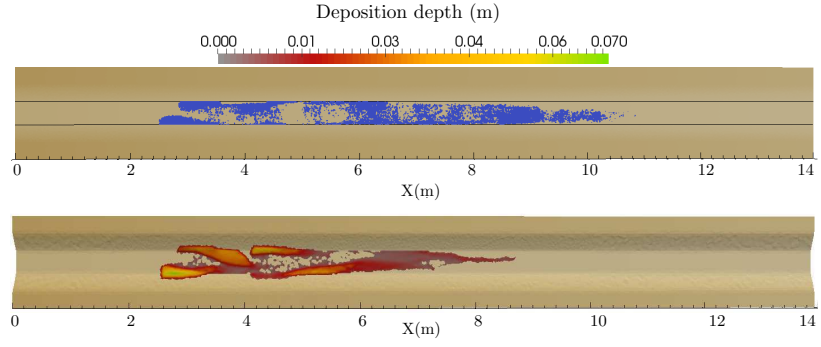
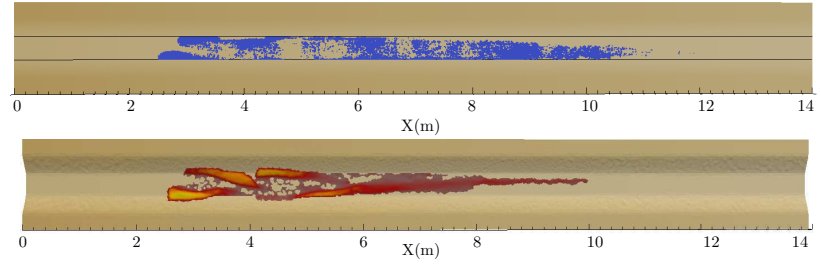


Figure 6: 2D experimental data of the covered surface (top views) and 3D computed contour views simulated numerically (bottom views) for the deposition depth along the non-erodible channel at times $t= 10, 30$ and 120 min for configuration 3

Time $t = 10 \text{ min}$



Time $t = 30 \text{ min}$



Time $t = 120 \text{ min}$

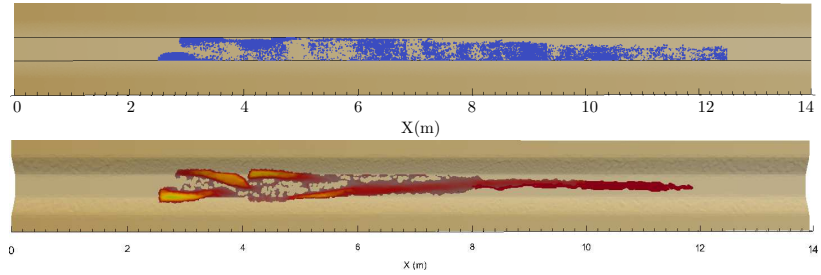
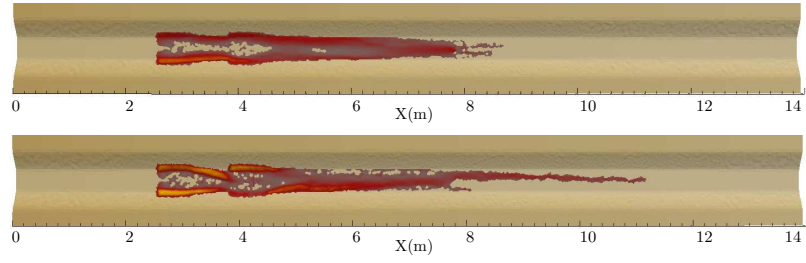


Figure 7: 2D experimental data of the covered surface (top views) and 3D computed contour views simulated numerically (bottom views) for the deposition depth along the non-erodible channel at times $t= 10, 30$ and 120 min for configuration 4

Configuration 3



Configuration 4

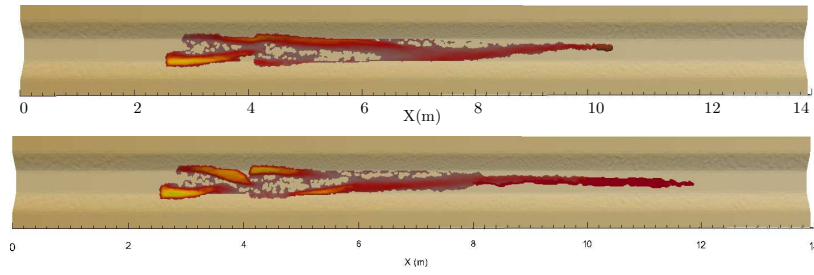
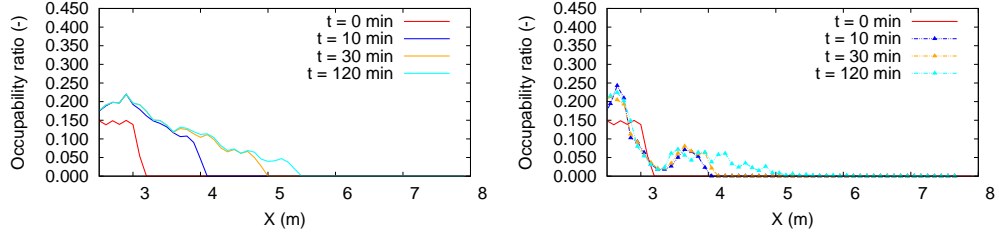
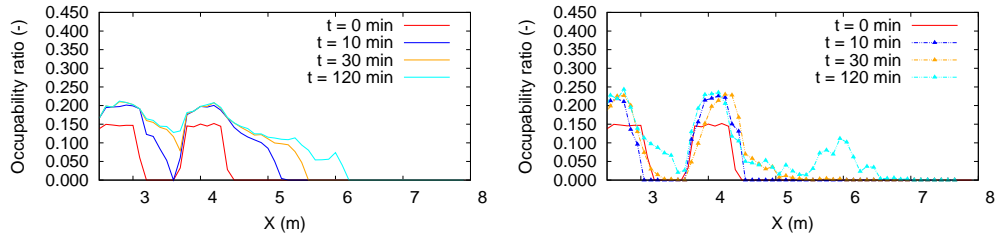


Figure 8: 3D computed contour views of the deposition depth for configurations 3 and 4 at time $t = 30 \text{ min}$ with a constant Manning parameter (upper) and with a dynamic Manning parameter (below)

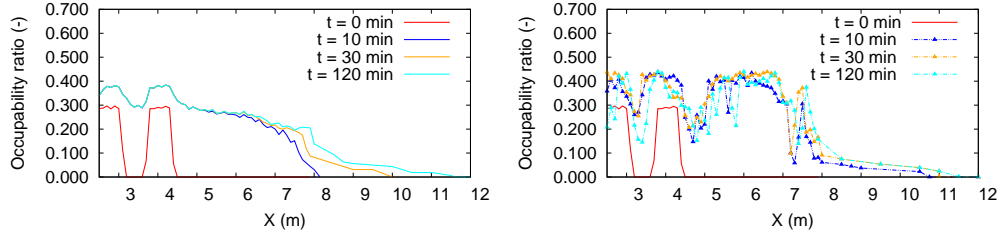
Configuration 1



Configuration 2



Configuration 3



Configuration 4

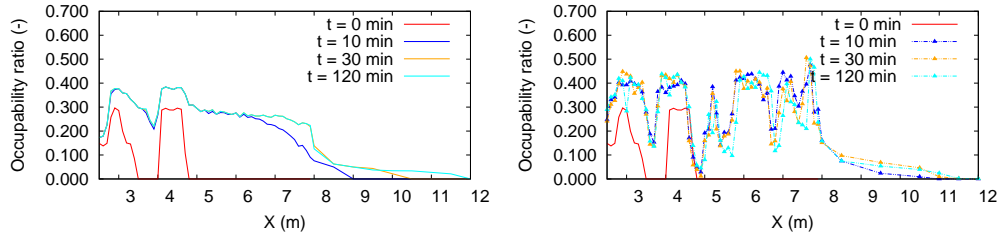


Figure 9: Temporal evolution of the occupability ratio (OR) for the four configurations: numerical results (left) and experimental data (right)

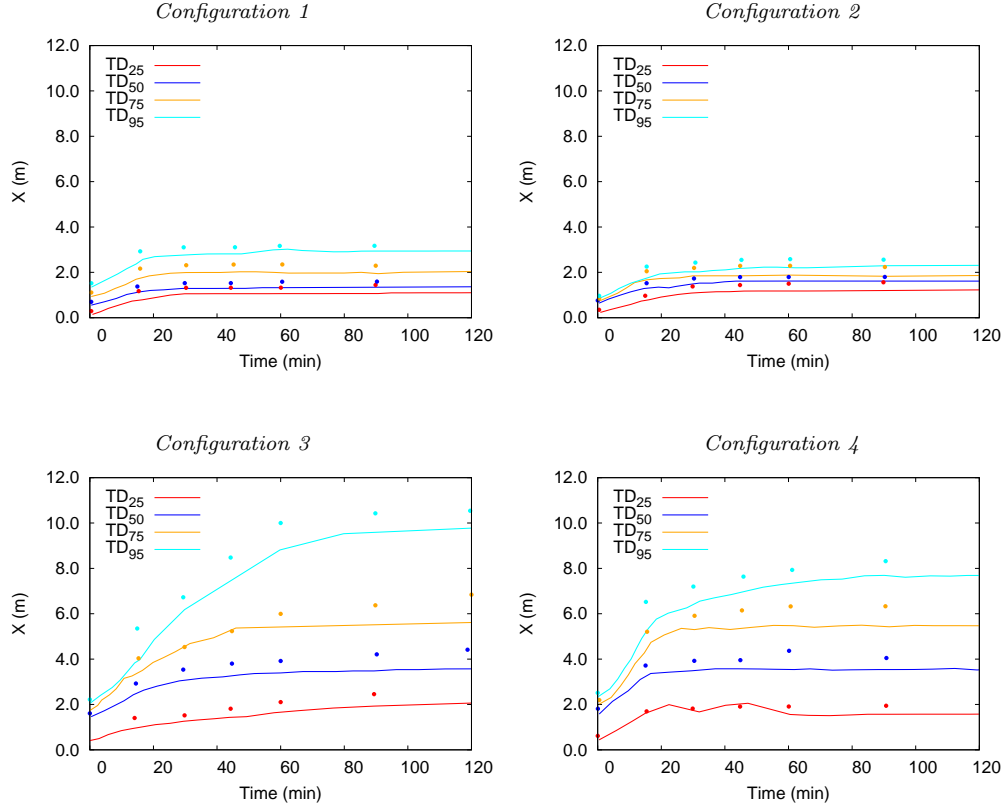


Figure 10: Temporal evolution of the travel distance (TD) for each configuration obtained with numerical simulation (continuous lines) and the experimental data (points)

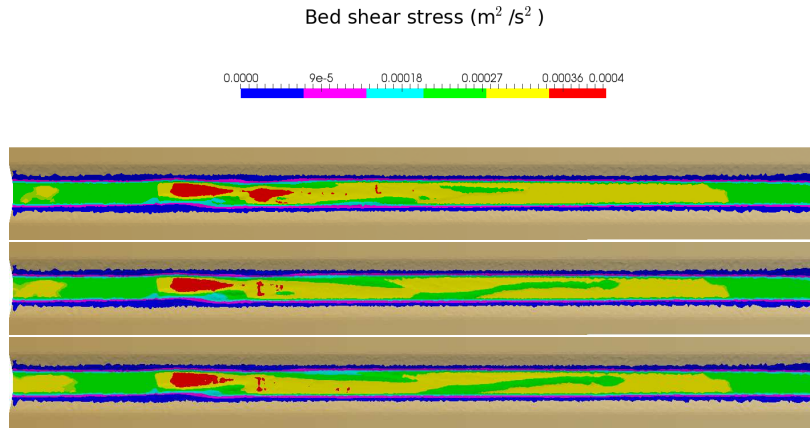


Figure 11: Computed contours of the bed shear stress at times $t = 10, 30$ and 120 min for configuration 4

5. Conclusions

Here, the potential of a numerical tool to assess the success of the application of replenishment of sediment in underfed armored channels is investigated by comparison with empirical data. The effect of several replenishment configurations in the downstream reach were simulated using a 2D Finite Volume scheme able to predict the sediment dynamics in combination with the flow behavior in a self-stable fashion. The mathematical model is built by means of the depth-averaged equations. In addition, the solid and liquid fluxes and source term discretization were derived from an approximate Riemann Problem. The forecasting abilities of the numerical tool when dealing with these environmental present real problems are herein showed and discussed.

Four different replenishment configurations have been studied. The first two configurations are simple but they allow to assess preliminarily the capabilities of the numerical tool. The deposition of the material in the downstream reach which is evaluated in terms of spreading using two parameters, occupational ratio and travel distance, is well tracked.

The third and fourth non-parallel configurations produce more complex deposition patterns and the interaction among blocks drive to a larger spreading. The patterns observed in laboratory are not fully predicted by the simulation. This is justified because in certain areas of the non-erodible channel there is a small number of particles whose transport behavior is mainly governed by collisional forces. The current numerical tool, based on a continuum approach, is still not suitable for this kind of particle-particle interaction. Nevertheless, the main replenishment structures are well captured in time and space by the simulation: maximum run-out, occupational ratio and travel distance.

Furthermore, another important remark is the temporal evolution of the bed roughness due to the new material supply which is provided by the replenishment. It is crucial to distinguish which characteristic grain diameter, between the non-erodible channel bed and the replenishment material, is dominating the bottom friction for obtaining accurate results. The methodology proposed in this work for updating continuously the friction factor leads to suitable results. The above findings show the usefulness of the considered model for the simulation of future different replenishment configurations.

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References

- [1] Brandt SA. Classification of geomorphological effects downstream of dams. *Catena* 2000;40(4):375–401.
- [2] Petts GE, Gurnell AM. Dams and geomorphology: Research progress and future directions. *Geomorphology* 2005;70(1-2):27–47.
- [3] Grant GE, Schmidt JC, Lewis SL. A Geological Framework for Interpreting Downstream Effects of Dams on Rivers. *Water Science and Application* 2005;7):209–225.
- [4] Power ME, Dietrich WE, Finlay JC. Dams and Downstream Aquatic Biodiversity: Potential Food Web Consequences of Hydrologic and Geomorphic Change. *Environmental Management* 1996;20(6):887–895.
- [5] Kondolf GM. HungryWater: Effects of Dams and Gravel Mining on River Channels. *Environmental Management* 1997;21(4):533–551.
- [6] Merz JE, Pasternack GB, Wheaton JM. Sediment budget for salmonid spawning habitat rehabilitation in a regulated river. *Geomorphology* 2006;76(1-2):207–228.
- [7] Kantoush SA, Sumi T, Suzuki T, Murasaki M. Impacts of sediment flushing on channel evolution and morphological processes: Case study of the Kurobe River, Japan. *Proc., River Flow 2010 International Conference on Fluvial Hydraulics*. 2010.
- [8] Balland P. Impacts des barrages sur les milieux physiques et biologiques. *Ingénieries: Supplément* 2004;38:23–32.
- [9] Ock G, Sumi T, Takemon Y. Sediment replenishment to downstream reaches below dams: implementation perspectives. *Hydrological Research Letters* 2013;7(3):54–59. doi:{doi.org/10.3178/hr1.7.54}.
- [10] Venditti JG, Dietrich WE, Nelson PA, Wydzga MA, Fadde J, Sklar L. Effect of sediment pulse grain size on sediment transport rates and bed mobility in gravel bed rivers. *Journal of Geophysical Research* 2010;115(3):1–19. doi:{[10.1029/2009jf001418](https://doi.org/10.1029/2009jf001418)}.
- [11] Zeh M, Donni W. Restoration of spawning grounds for trout and grayling in the river High-Rhine. *Aquatic science* 1994;56(1):59–69.

- 471 [12] Gaeuman D. Mitigating Downstream Effects of Dams. Gravel-Bed
472 Rivers: Processes, Tools, Environments 2012;Unique volume:182–189.
- 473 [13] Wheaton JM, Wheatonab JM, Pasternackc GB, Merz JE. Spawning
474 habitat rehabilitation – II. Using hypothesis development and testing
475 in design, Mokelumne River, California, U.S.A. International Journal
476 River Basin Management 2004;2(1):21–37.
- 477 [14] Gaeuman D. Recommended quantities and gradation for long-term
478 coarse sediment augmentation downstream from Lewiston Dam. 2008.
- 479 [15] Sklar LS, Fadde J, Venditti JG, Nelson P, Wydzga MA, Cui Y, et al.
480 Translation and dispersion of sediment pulses in flume experiments sim-
481 ulating gravel augmentation below dams. Water Resources Research
482 2009;45(8):n/a. doi:{10.1029/2008WR007346}.
- 483 [16] Lisle TE, Pizzuto JE, Ikeda H, Iseya F, Kodama Y. Evolution of a
484 sediment wave in an experimental channel. Water Resources Research
485 1997;8:1971–1981.
- 486 [17] Cui Y, Parker G, Pizzuto JE, Lisle TE. Sediment pulses in
487 mountain rivers: 1. Comparison between experiments and nu-
488 merical predictions. Water Resources Research 2003;39(9):1240.
489 doi:10.1029/2002WR001805.
- 490 [18] Ikeda H. Experiments on bedload transport, bed forms, and sedimentary
491 structures using fine gravel in the 4-meter-wide flume. Environmental
492 Research Center papers 1983;2:1–78.
- 493 [19] Venditti JG, Nelson P, Minear JT, Wooster J, Dietrich WE. Alternate
494 bar response to sediment supply termination. Journal of Geophysical
495 Research 2012;117(2):1–18. doi:{10.1029/2011jf002254}.
- 496 [20] Nelson PA, Brew AK, Morgan JA. Morphodynamic response of a
497 variable-width channel to changes in sediment supply. Water Resources
498 Research 2015;accepted:n/a. doi:{10.1002/2014WR016806}.
- 499 [21] Gaeuman D. High-Flow Gravel Injection for Constructing Designed in-
500 Channel Features. River Research and Applications 2014;30(6):685–706.
501 doi:{10.1002/rra.2662}.

- [22] Gomez B, Church M. An assessment of bed load sediment transport formulae for gravel bed rivers. *Water Resources Research* 1989;25(6):1161–1186. doi:{10.1029/WR025i006p01161}.
- [23] Aricò C, Tucciarelli T. Diffusive Modeling of Aggradation and Degradation in Artificial Channels. *Journal of Hydraulic Engineering* 2008;134(8):1079–1088. doi:{10.1061/(ASCE)0733-9429(2008)134:8(1079)}.
- [24] Wu W. Depth-Averaged Two-Dimensional Numerical Modeling of Unsteady Flow and Nonuniform Sediment Transport in Open Channels. *Journal of Hydraulic Engineering* 2004;130(10):1013–1024.
- [25] Soares-Fraza S, Zech Y. HLLC scheme with novel wave-speed estimators appropriate for two-dimensional shallow-water flow on erodible bed. *International Journal of Numerical Methods in Fluids* 2010;66(8):1019–1036. doi:{10.1080/00221686.2012.689682}.
- [26] Siviglia A, Stecca G, Vanzo D, Zolezzi G, Toro E, Tubino M. Numerical modelling of two-dimensional morphodynamics with applications to river bars and bifurcations. *Advances in Water Resources* 2013;52:243–260. doi:{dx.doi.org/10.1016/j.advwatres.2012.11.010}.
- [27] Canelas R, Murillo J, Ferreira RML. Two-dimensional depth-averaged modelling of dambreak flows over mobile beds. *Journal of Hydraulic Research* 2013;51(4):392–407. doi:{10.1080/00221686.2013.798891}.
- [28] Villaret C, Hervouet J, Kopmann R, Merkel U, Davies A. Morphodynamic modeling using the Telemac finite-element system. *Computers and Geosciences* 2013;53:105–113. doi:{10.1016/j.cageo.2011.10.004}.
- [29] Liu X, Landry B, García M. A hybrid Lattice Boltzmann method-finite difference method model for sediment transport and riverbed deformation. *River Research and Applications* 2015;31:447–456. doi:{10.1002/rra.2735}.
- [30] Juez C, Murillo J, García-Navarro P. A 2D weakly-coupled and efficient numerical model for transient shallow flow and movable bed. *Advances in Water Resources* 2014;71:93–109. doi:{10.1016/j.advwatres.2014.05.014}.

- 535 [31] Battisacco E, Franca MJ, Schleiss AJ. Physical modeling of artificial river replenishment techniques to restore morphological conditions
536 downstream of dams. Proceeding IAHR 2015 World Congress. ????
- 538 [32] Murillo J, García-Navarro P. Weak solutions for partial differential equations with source terms: Application to the shallow water equations. *Journal of Computational Physics* 2010;229:4327–4368.
540 doi:{10.1016/j.jcp.2010.02.016}.
- 542 [33] Juez C, Murillo J, García-Navarro P. Numerical assesment of bed load discharge formulations for transient flow in 1D and
543 2D situations. *Journal of Hydroinformatics* 2013;15(4):1234–1257.
544 doi:{10.2166/hydro.2013.153}.
- 546 [34] Smart G. Sediment transport formula for steep channels. *Journal of Hydraulic Engineering* 1984;3:267–276.
547
- 548 [35] Ferreira RML, Franca MJ, Leal JGAB, Cardoso AH. Flow over rough mobile beds: Friction factor and vertical distribution of the longitudinal mean velocity. *Water Resources Research* 2012;48:W05529.
549 doi:{10.1029/2011WR011126}.
- 552 [36] Ferreira RML. The von Kármán constant for flows over rough mobile beds. Lessons learned from dimensional analysis
553 and similarity. *Advances in Water Resources* 2015;81:19–32.
554 doi:{10.1016/j.advwatres.2014.10.004}.
- 556 [37] Manning R. On the flow of water in open channels and pipes. *Transactions of the Institution of Civil Engineers of Ireland* 1895;20:161–207.
557
- 558 [38] Strickler A. Beiträge zur Frage der Geschwindigkeitsformel und der Rauigkeitszahlen für Ströme, Kanäle und geschlossene Leitungen. *Mitteilungen des Eidgenössischen Amtes für Wasserwirtschaft* 1923;16 pp.
559
- 561 [39] Ferguson R, Hoey T, Wathen S, Werrity A. Field evidence for rapid downstream fining of river gravels through selective transport. *Geology* 1996;24(2):179–182.
562
- 564 [40] Hoey T, Bluck BJ. Identifying the controls over downstream fining of river gravels. *Journal of Sedimentary Research* 1999;69(1):40–50.
565

- 566 [41] Frings RM. Downstream fining in large sand-bed rivers. *Earth Science*
567 *Reviews* 2008;87:39–60.
- 568 [42] Wright S, Parker G. Modeling downstream fining in sand-bed rivers I:
569 Formulation. *Journal of Hydraulic Research* 2005;43(6):612–619.
- 570 [43] Wright S, Parker G. Modeling downstream fining in sand-bed rivers I:
571 Application. *Journal of Hydraulic Research* 2005;43(6):620–630.
- 572 [44] Manning M. Development of channel organization and roughness follow-
573 ing sediment pulses in single-thread, gravel bed rivers. *Water Resources*
574 *Research* 2001;37(8):2259–2272.
- 575 [45] Ferguson RI, Church M, Rennie CD, Venditti JG. Reconstructing a
576 sediment pulse: Modeling the effect of placer mining on Fraser River,
577 Canada. *Journal of Geophysical Research: Earth Surface* 2015;120:1436–
578 1454. doi:[10.1002/2015JF003491](https://doi.org/10.1002/2015JF003491).