

Holocene history of Aleppo pine (*Pinus halepensis* Mill.) woodlands in the Ebro Basin (NE Spain): Climate-biased or human-induced?

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Abstract

This paper reviews the past distribution of Aleppo pine woodlands in the Ebro Basin, Northeastern Iberia, from the Mesolithic to Modern times based on wood charcoal data. The aim is to detail the chronological timing and the drivers explaining the long-term presence of Aleppo pine woodlands and associated thermophilous flora. The available charcoal data support the early spread of *Pinus halepensis* during the Mesolithic (ca. 9000 cal BP) accompanied by Mediterranean trees and shrubs like *Quercus* sp. evergreen, *Juniperus* sp., *Arbutus unedo*, *Pistacia lentiscus*, *Rhamnus/Phillyrea*, Cistaceae, and *Rosmarinus officinalis*, as a local response to global climate change in the Early Holocene. During the arrival and the propagation of the Neolithic culture (ca. 7500–5500 cal BP), anthracological records, as well as regional palynological sequences, demonstrate the progressive replacement of a conifer-dominated open parkland by both *Quercus* sp. deciduous and evergreen woodlands in response to the Middle Holocene rise in temperature and humidity. This evidence, however, converges with the general idea that the presence and the spread of *Pinus halepensis* and associated scrubland have usually been attributed to the onset of landscape anthropization. The frequency of xero-thermophilous open scrubland and the use of Aleppo pine for fuel and wood crafting progressively increased during the Bronze and Iron Ages, and especially in Ibero-Roman and Medieval/Islamic times, when the vegetation landscape in the Middle Ebro Basin was largely deforested as a consequence of increasing demographic pressure, grazing and the establishment of proto-urban centers.

1. Introduction

Studies regarding the past distribution of key tree and shrub taxa at a European scale have increased noticeably during the last decades as a consequence of the effort of collecting and storing data in online databases by many paleobotanical research groups. The considerable number of available pollen and charcoal data that are chronologically well-constrained has allowed not only defining the location of both deciduous and evergreen taxa during the Last Glacial Maximum, but also characterizing the pattern of post-glacial recolonization of central and northern European regions (Brewer et al., 2002; Magri, 2008; Douda et al., 2014). Traditional biogeographical literature locates plant population refuge areas in the Mediterranean coastal areas (Bennett et al., 1991; Carrión et al., 2008), where warmer climate conditions prevailed during glacial environmental stages (Fletcher and Sánchez Goñi, 2008; Combourieu-Nebout et al., 2009). Only few studies have highlighted the role of inner continental enclaves as potential biodiversity reservoirs, despite the presence of mesic and thermophilous taxa has been widely recognized (González-Sampériz et al., 2004, 2010; Manzano et al., 2017).

One of the mostly widespread tree in semiarid Mediterranean Iberia is the Aleppo pine (*Pinus halepensis* Mill.), whose current distribution is confined to the thermo- and meso-Mediterranean biogeographic belts (Rivas Martínez, 1987). It is a post-fire obligate seeder, which does not survive fire but regenerates from seeds in post-fire environment (Ne'eman et al., 2004; Ne'eman and Izhaki, 2006; Pausas et al., 2004). For this reason, it is a common pioneer and invasive species within its natural distribution area (Pausas et al., 2002). Thus, the presence and the concomitant spread of *Pinus halepensis* and associated thermophilous scrubland have usually been attributed to the onset of landscape anthropization in the Mediterranean paleoecological literature (Thiébaud, 1997), without considering the role of this species as part of the natural vegetation. Besides, Costa (1987) and Braun-Blanquet and Bolós (1987), among others, define the stands of Aleppo pine as a secondary and marginal forests in Iberia, including the northeastern region belonging to the Ebro Basin. Indeed, the phytosociological literature discards the idea of self-established pine woodlands, in most cases associated with particularly unstable and/or immature topographic and soil conditions, cataloging them as “paraclimax” communities (Quézel and Barbero, 1990). Carrión and Fernández (2009) demonstrate that phytosociological theories do not always match with paleobotanical evidence, especially regarding Mediterranean pines and their presence and role since millennia in Iberia (Rodríguez-Ariza, 1992; Badal et al., 1994; Carrión et al., 2000; Rubiales et al., 2009; Olmedo-Cobo et al., 2019) and across the Mediterranean Basin (Vernet and Thiébaud, 1987; Wengler and Vernet, 1992). Therefore, the naturalness of Mediterranean pine populations in Iberia is still a matter of debate.

Similar to the presence of Aleppo pine, the progressive rise or even the long-term presence of xero-thermophilous scrubland has been widely taken as a clear sign of anthropogenic deforestation and landscape opening (Thiébaud, 1997). These communities, however, constitute an essential component of the Mediterranean *machia*, whose presence is linked to particular bioclimatic and ecological constraints (Blanco-Castro et al., 1997). Early-Middle Holocene pollen and charcoal data have shown the progressive replacement of open *Juniperus* parklands by *Olea–Pistacia–Quercus* sp. evergreen assemblages in the circum-Mediterranean region, interpreted as a direct

consequence of changes in precipitation and temperature regimes instead of human shaping (Wengler and Vernet, 1992; Tinner et al., 2009; Linstädter and Zielhofer, 2010; Di Rita and Melis, 2013; Djamali et al., 2013; Revelles et al., 2019). Thus, the distinction between human-made landscapes and climate-biased vegetation dynamics, the timing of appearance of the tree/shrub component and the forest disturbance/succession models, should be carefully interpreted, since they can produce similar pollen and charcoal assemblages (Figueiral and Mosbrugger, 2000).

The main goal of this paper is to outline the history of Aleppo pine woodlands and associated thermophilous flora in the Ebro Basin during the Holocene by reviewing published and unpublished charcoal assemblages. Therefore, we aim to discuss the possible drivers explaining the long-term dynamic of Aleppo pine communities in the Ebro Basin.

2. Regional setting

The Ebro River constitutes a huge basin, which runs from the Cantabrian Range to the Mediterranean across the northeastern Iberian Peninsula. It is bordered by the Pyrenees to the north, by the Iberian Range to the south and by the Catalanian Range to the east (Fig. 1A). It was formed by Alpine tectonic shifts that pushed up the outlying mountain chains and formed the sedimentary foreland depression. During the Tertiary, it was gradually filled with sediments as an endorheic continental basin. After the Late Miocene–Pliocene, as it became exoreic, it gained an outflow to the Mediterranean and the sediment moved to its bottom (200–300 m a.s.l.). This was the onset of the Ebro River system and of sediments draining from the basin to the sea, which still continues today.

A remarkable NW–SE climate gradient is shown across the Ebro Basin, determined by the strong Atlantic influence in the north and northwest during most of the year and the Mediterranean influence to the east, leaving a strong water-deficit during summer months. Three main sectors can be distinguished according to biophysical features: Upper Ebro Basin, Middle Ebro Basin and the Lower Ebro Basin (Fig. 1A).

The Upper Ebro Basin is climatically placed in the boundary between the Atlantic and the Mediterranean realm, which explains the fact that different regions have a deficit in summer rainfall. Overall, the vegetation is well-adapted to a mesic environment, mostly composed of *Quercus robur* and *Q. faginea* with *Acer campestre*, *A. monspessulanum*, *Fraxinus excelsior*, *Viburnum lantana*, and *Ligustrum vulgare* in the valley bottoms while higher elevations are dominated by *Quercus pyrenaica*, *Q. petraea*, *Betula pendula*, and *Fagus sylvatica*, depending on altitude, exposure and soil properties (Blanco-Castro et al., 1997). *Quercus ilex* develops on dry south-facing exposures along with *Phillyrea latifolia*, *Arbutus unedo*, *Viburnum tinus*, *Juniperus communis*, *Rhamnus alaternus*, *Hedera helix*, *Rubia peregrina*, and *Ruscus aculeatus*. The mountain ranges located in the Middle Ebro Basin reveal a marked altitudinal zonation. The basal belt is dominated by a Mediterranean woodland mainly composed by *Quercus ilex* and *Quercus faginea* communities, along with many xero-thermophilous shrubs such as *Quercus coccifera*, *Rosmarinus officinalis*, *Thymus vulgaris*, and *Genista scorpius*. *Pinus halepensis* is also present (Fig. 2) (Blanco-Castro et al., 1997). Higher altitudes support broadleaved elements like acidophilous *Quercus pyrenaica* and *Q. robur*, as well as *Sorbus aria*, *S.*

162 *aucuparia*, *Sambucus racemosa*, *Ilex aquifolium*, and *Fagus sylvatica*, among others
163 (Longares, 2004).

164 The lowlands of the Middle Ebro Basin form a steppe landscape, mostly dedicated to
165 irrigated agriculture. Annual rainfall values barely reach 350 mm (Fig. 1B).
166 Evapotranspiration rates are high (N800 mm) and temperature values are mostly related
167 to exposure and orientation (Fig. 1C). The vegetation cover is less than 50% and is
168 dominated by cereal crops with only small patches of open parkland of *Pinus halepensis*
169 (Fig. 2), *Quercus coccifera* and *Juniperus phoenicea*, and/or a dense scrubland with
170 *Rhamnus lycioides*, *Rosmarinus officinalis*, *Globularia alypum*, *Ephedra nebrodensis*, *E.*
171 *fragilis*, *Thymelaea tinctoria*, *Pistacia lentiscus*, *Phillyrea angustifolia*, and
172 *Brachypodium ramosum*, depending on topography and soil type. Mesophytes are
173 restricted to particularly humid mid-mountain enclaves (Blanco-Castro et al., 1997).
174 Nitrophilous and gypsophilous plants are abundant and include *Salsola vermiculata*,
175 *Atriplex halimus*, *Artemisia herba-alba*, *Peganum harmala*, *Ferula communis*, *Malcomia*
176 *africana*, *Marrubium alysoon*, *Ononis tridentata*, *Gypsophila hispanica*, *Helianthemum*
177 *squamatum*, and *Cistus clusii*.

178 Close to the Mediterranean, namely in the Lower Ebro Basin, the precipitation shows a
179 marked seasonality and is more abundant than in the inner sector (Fig. 1B). The mean
180 annual temperature reaches ca. 17 °C (Fig. 1C). The vegetation is well-adapted to meso-
181 and thermo-Mediterranean environments and is composed of evergreen oak forests with
182 small communities of deciduous oak occupying northern slopes and shaded areas.
183 Towards the coast, a thermophilous community is found in where *Quercus coccifera*,
184 *Olea europaea*, *Pistacia lentiscus*, *Pinus halepensis*, and *Chamaerops humilis*, among
185 others, acquires special relevance (Folch, 1986).

186 The Ebro Basin has been intensely populated by prehistoric human groups since the
187 Paleolithic. The character of the Ebro Basin as crossroad for people and fauna makes it
188 an area with frequent evidence of long term human occupation that makes it a territory
189 with a particular cultural character (Aura et al., 2011; Utrilla et al., 2012; Domingo et al.,
190 2018). The first human impact on the environment is documented in Neolithic times
191 (Montes et al., 2016; Alcolea et al., 2017a) and progressively increased in recent times,
192 paralleling the rise of human peopling and social complexity.

193 3. Material and methods

194 The anthracological data reviewed for the Ebro Basin come from different sedimentary
195 archives, including archeological sites and burial paleosols (Fig. 1A and Table 1). A
196 number of 27 sites reporting *Pinus halepensis* wood charcoal have been considered for
197 the present work, which comprises chronologies ranging from the Mesolithic until the
198 Modern Ages.

199 The taxonomic distinction between Mediterranean and cryophilous montane *Pinus*
200 species, based on anatomical diagnostic traits, is usually feasible when allowed by a good
201 state of preservation of wood remains. The wood of cryophilous montane European pines
202 (*Pinus nigra* ssp. *salzmanni* (Dunal) Franco, *Pinus sylvestris* L. and *Pinus mugo* Turra)
203 is commonly grouped under the taxon *Pinus sylvestris* type, *Pinus nigra/sylvestris* type,
204 *Pinus nigra/sylvestris* or *Pinus sylvestris/uncinate* (Figueiral and Carcaillet, 2005; Roiron

et al., 2013; Allué et al., 2018). Mediterranean pines (*Pinus halepensis* Mill., *Pinus pinaster* Ait., *Pinus pinea* L.) can frequently be determined at the species level.

Pinus halepensis is a softwood characterized in transverse section by gradual early-latewood transition and the presence of resin canals throughout the growth ring. They are generally smaller than in other species of the genus (Fig. 3). In the longitudinal sections, the average height of rays is 10 cells, the upper limit being 22 cells. The most characteristic anatomic features are ray parenchyma cells with 1–3 or 4 pinoid pits in cross-fields and the presence of slightly dentate walls in ray tracheids (Jacquot, 1955; Schweingruber, 1990) (Fig. 3).

In order to facilitate readability and infer landscape changes across different cultural periods, some charcoal data have been grouped according to ecological affinities (Fig. 4).

- Xero-thermophytes comprises taxa related to open scrubland and the Mediterranean machia, namely, *Arbutus unedo*, *Buxus sempervirens*, cf. *Berberis* sp., *Chamaerops humilis*, Chenopodiaceae, Chenopodiaceae cf. *salsola*, Cistaceae, *Cistus* sp., *Erica* sp., Ericaceae, *Ephedra* sp., Fabaceae, Labiatae, Labiatae cf. *Lavandula*, Monocotyledoneae, *Olea europaea*, cf. *Olea europaea*, *Rhamnus/Phillyrea*, *Pistacia* sp., cf. *Pistacia* sp., *Pistacia lentiscus*, *Pistacia* cf. *terebinthus*, *Rosmarinus officinalis*, and *Retama monosperma*.

- Mesophytes and riparian taxa include *Abies alba*, *Acer* sp., *Alnus* sp., *Betula* sp., *Celtis australis*, cf. *Celtis*, *Corylus avellana*, *Corylus-Alnus*, *Fagus sylvatica*, *Fraxinus* sp., *Populus* sp., *Populus/Salix*, Rosaceae Maloideae, cf. Rosaceae Maloideae, Rosaceae Pomoideae, Rosaceae Prunoideae, Rosoideae, *Prunus avium/spinosa* type, *Salix* sp., Salicaceae, *Sorbus* sp., *Tamarix* sp., *Taxus baccata*, *Ulmus* sp., *Viscum* sp., and *Vitis vinifera*.

- *Pinus* sp. cryophilous comprises *Pinus nigra/sylvestris* type and *Pinus sylvestris* type.

- *Pinus* sp. other is composed of *Pinus* sp., *Pinus* bract, *Pinus* cortex, *Pinus halepensis/pinaster* type, *Pinus pinea/halepensis* type, and *Pinus pinea/pinaster* type.

- The group of other taxa includes species with generic ecological affinities and poor descriptions, namely, Angiosperm indeterminate, Conifer indeterminate, *Prunus* sp., and *Quercus* sp.

- Cultivated taxa group is represented by *Ficus carica*, *Juglans regia*, *Prunus amygdalus* type, *Prunus dulcis*, cf. *Punica granatum*, and *Pyrus malus*.

These reviewed wood charcoal assemblages were chronologically constrained by multiple ¹⁴C AMS dates submitted to different laboratories. Charcoal-based radiocarbon dates were calibrated using CALIB 7.1 software (Stuiver and Reimer, 1993) based on the latest IntCal 13 curve (Reimer et al., 2013). The majority of radiocarbon dates are derived from sedimentary deposits although direct dating of charred woody elements from *Pinus halepensis* were also obtained (Table 3).

4. Results

The collection of wood charcoal data in the Ebro Basin has been strongly linked to archeological contexts. Anthracological analysis has largely increased since the beginning of the 21th century with the progressive intensification of paleobotanical studies and the publication of many reviews centered in the Upper (Ruíz-Alonso, 2014; Ruíz-Alonso and Zapata, 2015; Zapata, 2000; Zapata and Peña Chocarro, 2005) and Middle-Lower Basin (Carrión, 2007; Piqué et al., 2012; Vila and Piqué, 2012; Alcolea 2017, 2018; Vila, 2018). Nevertheless, a gap of knowledge persists, depending on regions and periods.

Upper Pleistocene wood charcoal records (e.g., Cova Gran, Cueva del Gato) show the predominance of montane pinewoods and junipers revealing an open landscape with scattered conifers, all of them well adapted to cold (Piqué, 1995; Allué et al., 2018; Mazo and Alcolea, 2019) and especially dry climate conditions in the valleys during the Last Glacial Maximum (Badal et al., 2012). The Late Glacial (ca. 13,500 cal BP) charcoal spectra reveal a progressive spread of mesic taxa (mainly *Quercus* sp., *Acer* sp., *Prunus* sp., and *Corylus avellana*) although conifers continue being the main landscape components (Allué et al., 2013; Montes et al., 2016). To our knowledge, no Pleistocene Aleppo pine charcoal remains have been reported for the Ebro Basin.

During the onset of the Holocene, local variations in vegetation landscapes are inferred from anthracological assemblages placed between the high and middle mountain regions and the Central Depression. The Pre-Pyrenean Mesolithic records (Allué et al., 2012; García-Martínez de Lagrán et al., 2016; Montes et al., 2016; Ruíz-Alonso and Zapata, 2015) are dominated by montane pine communities with scattered presence of broadleaved elements like *Quercus* sp. deciduous, *Acer* sp., *Fraxinus* sp., *Ulmus* sp., Rosaceae Maloideae at ca. 11,350–8000 cal BP, fitting the typical post-glacial forest succession of the Eurosiberian realm (Carrión et al., 2010; Iriarte-Chiapusso et al., 2016). By contrast, *Pinus halepensis* is especially relevant in the lowlands accompanied by Mediterranean trees and shrubs like *Quercus* sp. evergreen, *Juniperus* sp., *Arbutus unedo*, *Pistacia lentiscus*, *Rhamnus/Phillyrea*, *Prunus* sp., Cistaceae, and *Rosmarinus officinalis* at ca. 9000 cal BP, as revealed by Mesolithic sites, such as Cabezo de la Cruz (Badal, 2013), Espantalobos (Alcolea et al., 2017b), El Pontet (Alcolea, 2017b) and Los Baños de Ariño (Badal, 2004), as well as the nearby alluvial sequence of La Poza (Pérez-Lambán et al., 2018) (Fig. 1A and Table 2). Direct radiocarbon dates on wood charcoal and cone bracts support the local presence of pine communities around 8995–8625 cal BP (Table 3), revealing widespread Aleppo pine populations and associated flora in the Middle Ebro Basin during the Early Holocene.

During the Neolithic (7500–5500 cal BP), the previous open thermophilous landscape persisted, although mesophytes, and especially *Quercus* sp. evergreen, were more in the anthracological assemblages (Fig. 4 and Table 4). *Pinus halepensis* is documented in sites of the Bajo Aragón (Fig. 1A), namely Plano del Pulido (Alcolea, 2017b) and in the Lower Ebro Basin, such as the Ebro Delta (e.g., Barranc d'en Fabra, Ros, 1996). It is found in mid mountain areas in the eastern Pre-Pyrenees (e.g., Cueva de Chaves, Auvelles, Martín-Seijo and Piqué, 2008; Alcolea et al., 2017a) and nearby coastal enclaves like La Cativera, Cova Bonica or El Cavet (Fontanals et al., 2008; Allué et al., 2017). However, in the

Upper Ebro Basin it is not found together with juniper and evergreen oaks, as revealed by the study in the open-air site of Los Cascajos (Peña-Chocarro et al., 2005).

Bronze and Iron Age are well-known periods from the archeobotanical point of view, both for the good conservation of the remains compared to older times and for the large number of studies concentrated in La Plana de Lleida at the eastern Middle Ebro Basin (Fig. 1A). The vegetation landscape is dominated by Aleppo pine communities accompanied by *Juniperus* sp. and *Quercus* sp. evergreen. This is particularly evident in Minferri (Piqué et al., 2012; Vila and Piqué, 2012) and Pla de Tabac I (Esteve et al., 2015) during the Bronze Age but also in the so-called Segre-Cinca I (Punta Farisa, El Vilot 0, Roques del Sarró) and Segre-Cinca II (Genó, Masada de Ratón, Tossal de Solibernat, Vincamet I) archeological groups (Piqué et al., 2012; Vila and Piqué, 2012; Vila, 2018), as well as in Cabezo de la Cruz (Badal, 2009; Carrión, 2009), where the highest representation of *Pinus halepensis* is attained (Fig. 4 and Table 5). However, the frequency of Mediterranean open scrubland progressively acquired more visibility during the following cultural periods (Fig. 4). The Segre-Cinca III group (Vilot I, II and Vincamet II), Cabezo de la Cruz during the Iron Age as well as Iron Age settlements of Els Vilars (phases 0, I), La Codera, Tozal de los Regallos or Segeda (Fig. 4 and Table 5) reported a remarkable rise in xero-thermophilous scrubland with *Pistacia lentiscus*, *Rosmarinus officinalis*, *Rhamnus/Phillyrea*, and Cistaceae, denoting a progressive landscape opening (Carrión, 2007; Badal, 2009; Piqué et al., 2012; Vila and Piqué, 2012; Vila, 2018). The wood charcoal assemblages of the Iberian period reveal a trend towards an open heliophyte scrubland (e.g., Estinclells, Missatges, Roques del Sarró; Piqué et al., 2012; Vila and Piqué, 2012) (Table 6). The increase of both *Quercus* sp. Deciduous and *Quercus* sp. evergreen during this phase (Fig. 4) could be related to the depletion of resources close to the archeological site and their search in remote territories. In this case these taxa would not correspond to the local vegetation.

Roman, Medieval/Islamic and Modern times are poorly known due to the almost complete absence of wood charcoal studies from archeological sites. Regional paleoenvironmental studies reveal a sharp landscape transformation in the Middle Ebro Basin (Pérez-Lambán et al., 2018), progressively shaping the current vegetation distribution. Aleppo pine is one of the most exploited taxa for fuel in Roman Ilerda and, to a lesser extent, in Iesso (Fig. 4A), being accompanied by *Pinus* type *nigra/sylvestris*, *Quercus* sp. deciduous, *Quercus* sp. Evergreen and Fabaceae, among others (Piqué et al., 2012) (Table 7). The charcoal spectra evidenced the spread of arboricultural practices across the basin during the Medieval/Islamic period, incorporating tree elements of great economic value, such as *Juglans regia*, *Prunus spinosa/avium* type, *Prunus amygdalus* type, *Pyrus malus*, and Rosaceae Maloideae to local economies (Alonso et al., 2014; Alcolea et al., 2016) (Table 7). *Pinus halepensis* continued being exploited as an important economic resource in Madíña Làrida, Pla d'Almatà and Juslibol Castle (Piqué et al., 2012; Vila and Piqué, 2012; Alcolea et al., 2016, respectively), while it was the dominant taxon in the pedoanthracological record of Val de Zaragoza (Fig. 1A), chronologically framed within the Little Ice Age (Peña-Monné et al., 2018) (Fig. 4 and Table 7). Agricultural terraces and buried horizons date around the 18th century. They reveal an open scrubland where many thermophytes (e.g., *Rosmarinus officinalis*, *Pistacia lentiscus*) were present (Boixadera et al., 2016).

5. Discussion

The results of our survey reveal that the modern distribution of *Pinus halepensis* in the Ebro Basin is not only human-induced, but also the consequence of a combination of factors, including long-term ecological responses to climate changes, soil character, and micro-topographic features. Overall, the reviewed paleobotanical data reveal woodlands of Aleppo pine in the Ebro Basin during the Mesolithic and demonstrate that its distribution is natural before the Neolithic landscape transformation.

The scarcity of wood charcoal data chronologically framed in the Ebro Basin and comprising Upper Pleistocene chronologies makes it difficult to evaluate the spatial extent of Aleppo pine communities since most of the known archeological sites are from mountainous areas. Up to now, anthracological data suggest the dominance of cryophilous pines (*Pinus sylvestris* type) from the NE-E Iberian Mediterranean coast to the Pyrenees (López-García et al., 2012; Allué et al., 2018). Some charcoal assemblages have shown the local presence of Mediterranean pines in SE Iberian sites in coastal areas during the Middle and Upper Paleolithic. The Mousterian and Solutrean charcoal assemblages from Cueva de Nerja (Badal, 1998), Cueva Antón (Zilhão et al., 2016), and Cueva de Ambrosio (Rodríguez-Ariza, 2006) document the presence of a thermophilous flora in which *Pinus halepensis* acquires special relevance. In any case, the Pleistocene presence of a thermophilous *machia* is restricted to the warmer areas of southern Iberia (Carrión et al., 2008; Badal et al., 2012).

This arises the question if these continental regions may have acted as refuge zones for the thermophilous flora, despite the lack of direct paleobotanical evidence comprising Last Glacial Maximum/Lateglacial chronologies. Long-term pollen records from the Middle Ebro Basin, such as Lake Estanya, support the expansion of conifer forests during the Lateglacial interstadial GI-1, when pine and juniper communities account N60% of total terrestrial pollen (González-Sampériz et al., 2017). Although the distinction among pines based exclusively on pollen morphology is not an easy task (Desprat et al., 2015), the coeval presence of pines and a meso-thermophilous flora suggests that inner continental enclaves may have constituted reservoirs of biodiversity that could expand regionally once warmer temperatures and moister conditions took place (González-Sampériz et al., 2004, 2010). In fact, the Aleppo pine communities placed in NE Iberia reveal a high genetic diversity in comparison to southerly placed populations (Gómez et al., 2005). This would imply a rapid migration from coastal enclaves to inner continental environments.

The onset of the Holocene brought a major change in terms of forest dynamics in the Iberian Peninsula, showing a lagged response of Mediterranean ecosystems to the increase in humidity and temperatures in the Pre-Pyrenean region and the Central Depression (Morellón et al., 2018). In the Early Holocene, climate model simulations highlight an enhanced seasonality in air temperatures and occurrence of dry summer conditions in the Mediterranean region (Brayshaw et al., 2011), which may explain reduced lake-levels (Morellón et al., 2009; Aranbarri et al., 2014) and limited fluvial activity in the Ebro Basin and the Iberian Range (González-Amuchastegui and Serrano, 2015; Sancho et al., 2015; Aranbarri et al., 2016).

377 These continental climate constraints may have supported a conifer dominated landscape
378 in the Ebro Basin until ca. 9500 cal BP (Morellón et al., 2018). Long-term pollen records
379 reveal open woodlands, almost exclusively dominated by pines and junipers, which were
380 progressively replaced by evergreen oaks and thermophilous scrubland towards the
381 Middle Holocene (Aranbarri et al., 2014; González-Sampériz et al., 2017).
382 Anthracological data show that continental *machia* was already present in the lowlands
383 as revealed by all reviewed Mesolithic settlements (Fig. 4), with Aleppo pine woodlands
384 as important landscape components (Fig. 5), amply used as firewood in Cabezo de la Cruz
385 and Espantalobos sites (Badal, 2013; Alcolea et al., 2017b) (Fig. 1A and Table 2). At that
386 time (8340–8100 cal BP), Aleppo pine was also attested in the Levantine coastal areas,
387 such as in La Falguera (Carrión et al., 2006) or Ebro Basin borderlands as La Cativera
388 (Allué et al., 2000). Its presence was also reported from other regions of inland Iberia,
389 such as Cueva Blanca, where charcoal spectra define an open Mediterranean landscape
390 with Aleppo pine communities, junipers and thorny scrubland between 8400 and 7500
391 cal BP (Uzquiano et al., 2016). The accompanying flora in the Mediterranean coast,
392 however, clearly differs from the continental realm, where humidity-demanding species
393 are less represented (Fig. 4 and Table 2). In any case, the presence of Aleppo pine prior
394 to the Neolithic is documented in many Iberian archeological sites (Badal and Roiron,
395 1995; Allué, 2002; Aura et al., 2005; Alcolea, 2017a, among others).

396 The review of wood charcoal assemblages and the palynological evidence of lacustrine
397 paleoenvironmental records suggest that the regional presence of Mesolithic woodlands
398 in the Ebro Basin was driven by a particular continental climate, as it occurs nowadays,
399 rather than being direct consequence of human activities. Hunter-gatherer societies may
400 have modified forests composition at a local-scale by obtaining resources for fuel and
401 timber wood, but this cannot explain the long-term persistence of Aleppo pine woodland
402 in continental Iberia during the first millennia of the Holocene (Table 2). Thus, the
403 climatic character of this community in the Ebro Basin should be addressed.

404 The scarcity of continuous archeological pollen records comprising the
405 Mesolithic/Neolithic transition make it difficult to detect the early impact of farming
406 practices in the Ebro Basin. Thus, there has been much discussion about the timing and
407 modes of landscape colonization (Oms et al., 2016; Alday et al., 2018; Rojo-Guerra et al.,
408 2018; Laborda, 2019). The available pollen records reveal minor shifts in the arboreal
409 pollen values, with punctual peaks in the herbaceous components and/or scattered
410 presence of cereal pollen (e.g., Cueva de Chaves: López García and López Sáez, 2000).
411 These assemblages clearly differ from emblematic Neolithic settlements with high
412 demographic pressure and intense landscape transformations, such as La Draga (Revelles
413 et al., 2015), Conquezueta-Ambrona Valley (Aranbarri et al., 2015), and El Mirador
414 (Expósito et al., 2017). Thereby, the common idea that *Pinus halepensis* woodlands
415 spread as secondary forests linked to human activities is not supported by archeological
416 pollen profiles, continuous lacustrine pollen records and wood charcoal assemblages
417 obtained in the Middle Ebro Basin (Fig. 4 and Table 4). In fact, arboreal pollen
418 percentages attained their highest values in continental Iberia between 9000 and 6000 cal
419 BP, in response to the establishment of humid conditions and warmer temperatures
420 (Carrión et al., 2010; Morellón et al., 2018).

The hypothesis that human pressure was a major driver, able to completely change the local vegetation dynamic and woodland composition, is not highlighted by the reviewed sites. Moreover, charcoal assemblages (Fig. 4) show that Aleppo pinewoods are less represented in the Ebro Basin during the Neolithic (7500–5500 cal BP) than the Mesolithic times (9000–7500 cal BP). If Neolithic economic activities and landscape transformations were the main factor favoring the spread of Aleppo pine as secondary forest, its frequencies in the paleobotanical results would be expected to increase in all the sites (Fig. 4 and Table 4). This pattern is also confirmed in coastal Mediterranean sequences such as Cova Bonica in where *Pinus halepensis* values are highly represented in the Early Neolithic (7400–7200 cal BP) but decreases until disappearance in the Late Neolithic (4800–3400 cal BP) when an open scrubland composed of *Erica*, *Arbutus unedo* and *Quercus* sp. Evergreen dominates the landscape (Allué et al., 2017). Charcoal assemblages indicate the dominance of both deciduous and evergreen oaks (Fig. 5), which were among the main exploited taxa in Neolithic times, as confirmed by pollen results obtained from lacustrine records and tufa build-ups at a regional scale (Aranbarri et al., 2014, 2016; González-Sampériz et al., 2017; Luzón et al., 2017). This evidence reinforces the hypothesis that the vegetation change was driven more by a climate forcing than by an indiscriminate use of anthropogenic fire for clearing purposes. In the same time period, independent paleohydrological (Morellón et al., 2009; Aranbarri et al., 2014) and geomorphological proxies (Aranbarri et al., 2016; Luzón et al., 2017) document a climate shift towards moister conditions in the Ebro Basin. Thus, a more mesic climate may partially explain the reduction of pinewoods suggested by the paleobotanical assemblages, as Aleppo pine is physiologically better adapted to dry and arid conditions (Borghetti et al., 1998) and is less competitive against both deciduous and evergreen oaks under humid climate constrains (Zavala et al., 2000). These results are not incompatible with a local shaping of the vegetation landscape documented in many archeological sites, but force us to reconsider that human activities were the major factor favoring the spread of pine populations at a regional scale, at least during the Early-Middle Holocene transition.

The Late Holocene phase in the Ebro Basin is defined as a period of rapid and intense environmental changes (Morellón et al., 2009; Pérez-Lambán et al., 2014). Pollen profiles reveal a long-term substitution of mixed Mediterranean woodland by an open conifer forest dominated by pine woodlands (González-Sampériz et al., 2017). Broadleaved elements almost disappear or are confined to humid enclaves near watercourses. Similarly, independent proxies, such as the geochemical signature of lacustrine sediments and geomorphic evidence related to tufa accumulation, reveal a major climate change towards reduced water availability at a basin scale (Morellón et al., 2009; Aranbarri et al., 2016; Luzón et al., 2017).

Anthracological data reveal a progressive reduction of woodlands and the establishment of a Mediterranean scrubland dominated by *Arbutus unedo*, *Rhamnus/Phillyrea*, *Rosmarinus officinalis*, *Pistacia lentiscus*, and Fabaceae, chronologically framed during the Bronze and Iron Ages (Fig. 4 and Tables 5 and 6). Aleppo pine populations could have been partially favored by the establishment of a warmer and drier climate in the lowlands and mountain foothills similar to present day.

However, the increased demographic pressure documented during this period along the Ebro Valley (Blanco-González et al., 2018; Fyfe et al., 2019) necessarily implies a more frequent use of wood as fuel and timber (Carrión, 2007; Badal et al., 2010). The development of craft activities, namely pottery and metallurgy, demanded the use of fuel beyond the domestic sphere. The characteristics of Aleppo pine trees, a fast-growing species producing long and straight trunks, makes it widely used as timber wood (Ntinou et al., 2013; Picornell-Gelabert and Dufraisse, 2018). Although its use as raw material is difficult to document due to the scarcity of non-charred plant-based materials in archeological records, Aleppo pine woodcraft is attested in the Middle Ebro Basin (Alcolea and Rodanés, 2019) and in other Mediterranean areas (Carrión and Rosser, 2010) during Iberian and Roman times.

Differently from the Mesolithic (Table 3) and Neolithic (Table 4) assemblages, the Mediterranean scrubland component increased since the Metal Ages (Tables 5 and 6), and especially during historical times (Table 7), which may be interpreted as the result of a more intensive land use linked to higher demographic pressure and diversification of economic activities. Indeed, pollen data reveal a progressive spread of grazing activities, rise of cereal cultivation and introduction of arboricultural practices (Riera et al., 2004; Corella et al., 2013; González-Sampériz et al., 2017), witnessing a profound woodland depletion at basin scale and favoring the expansion of Aleppo pine communities in recent times.

6. Conclusions

The vegetation record of the Ebro Basin reveals a long-term dominance of Aleppo pine forest and Mediterranean scrubland since the Mesolithic, with the exception of the Neolithic, when a more humid demanding flora spread in the region. The degree of anthropogenic pressure across the Holocene, however, did not have always the same intensity, so the interaction between people and trees and the interplay with the regional climate background must be interpreted in its historical context.

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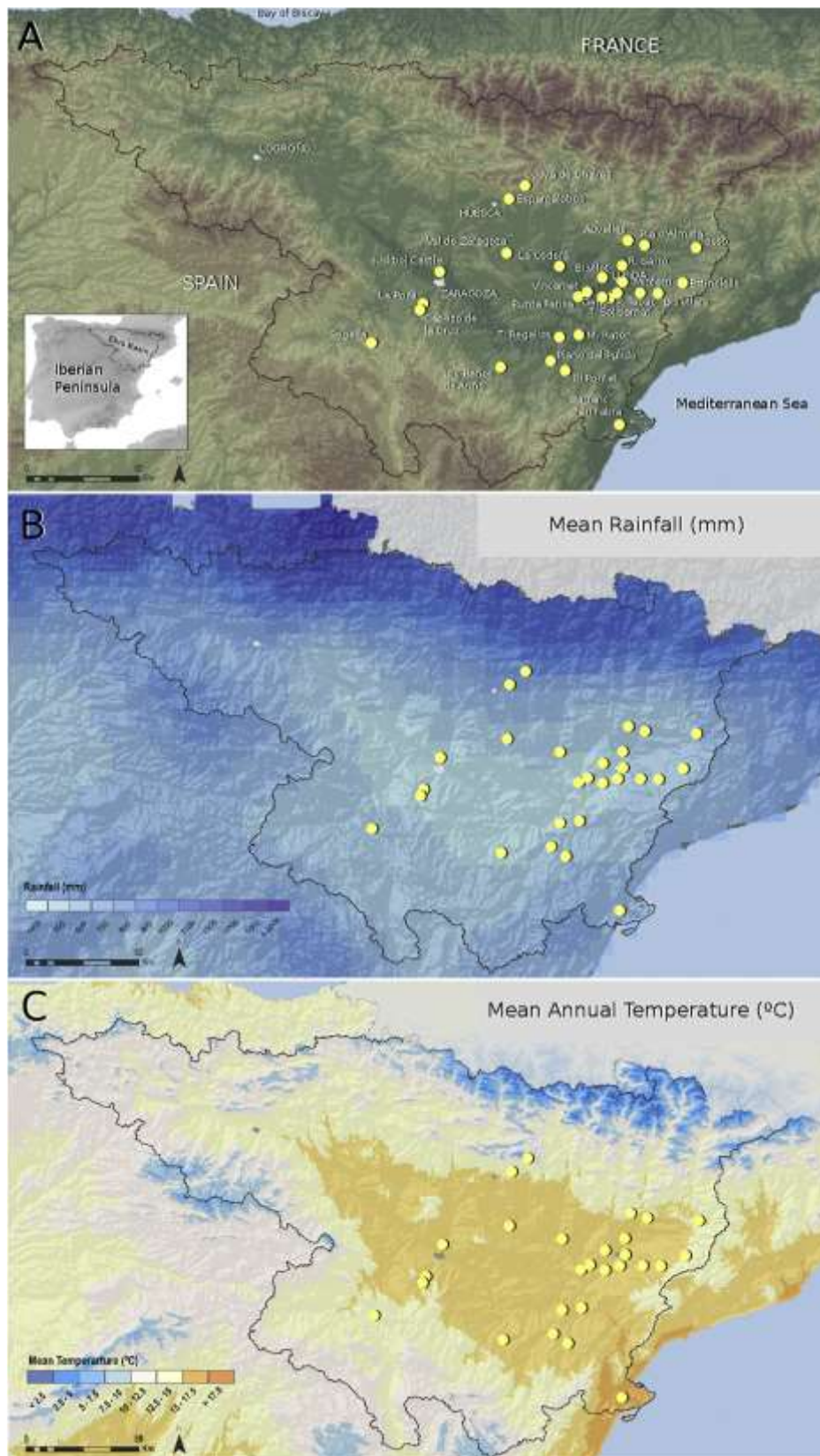


Fig. 1. A) Location map of the Ebro Basin reporting *Pinus halepensis* wood charcoal. Source: Confederación Hidrográfica del Ebro (<http://www.chebro.es>). B) Main rainfall features of the Ebro Basin and borderlands (González-Hidalgo et al., 2011). C) Mean annual temperature of the Ebro Basin and borderlands (González-Hidalgo et al., 2015).

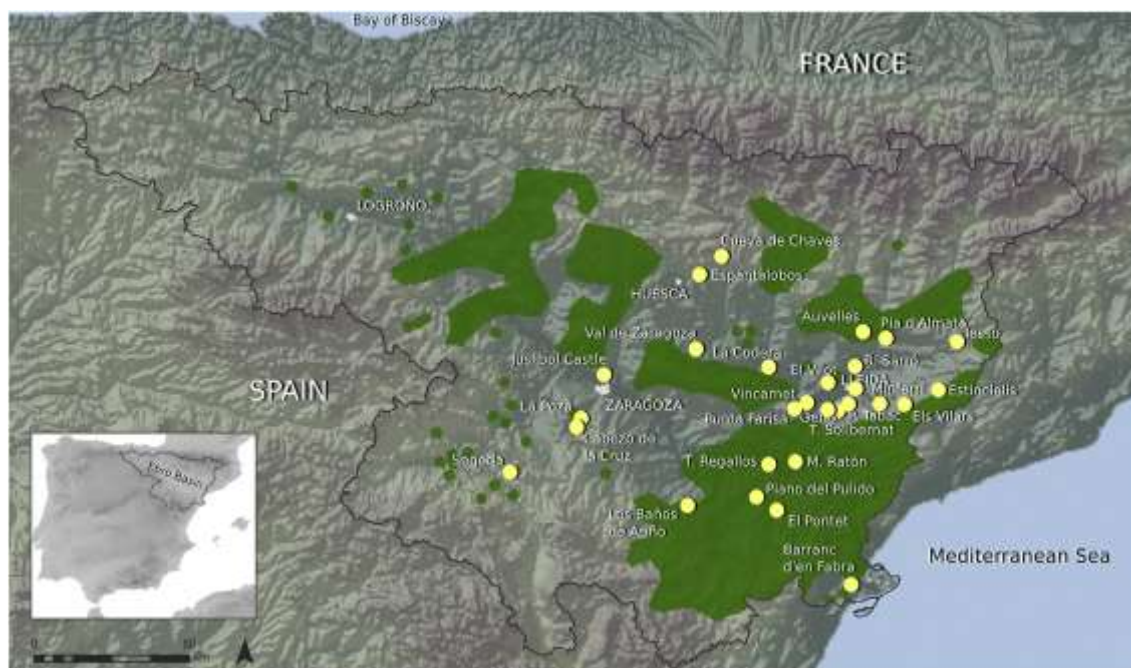


Fig. 2. Modern distribution of *Pinus halepensis* in the Ebro Basin. The references of the sites are shown in Table 1. Source: EUFORGEN (<http://www.euforgen.org/>).

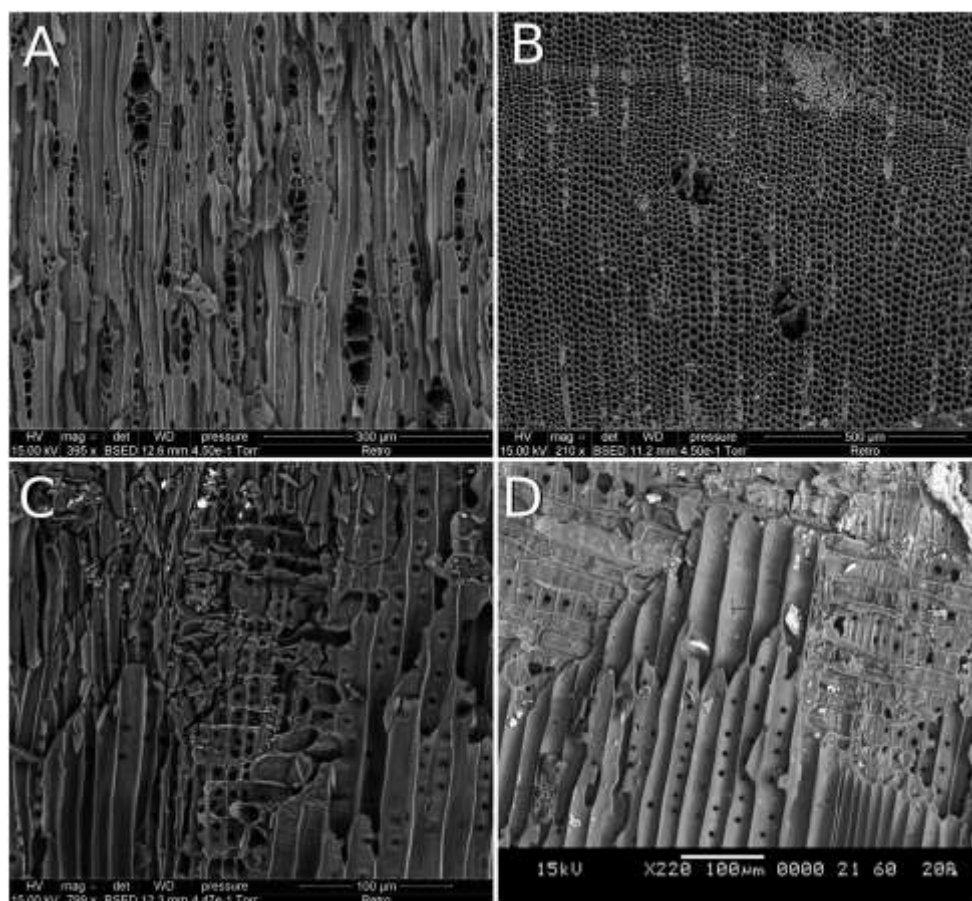


Fig. 3. Scanning Electron Microscope images of selected fragments of *Pinus halepensis*. Radial, transverse and tangential sections from Espantalobos archeological site (A–C, respectively) (Alcolea, 2018). D) Tangential section from El Pontet archeological site (Alcolea, unpublished).

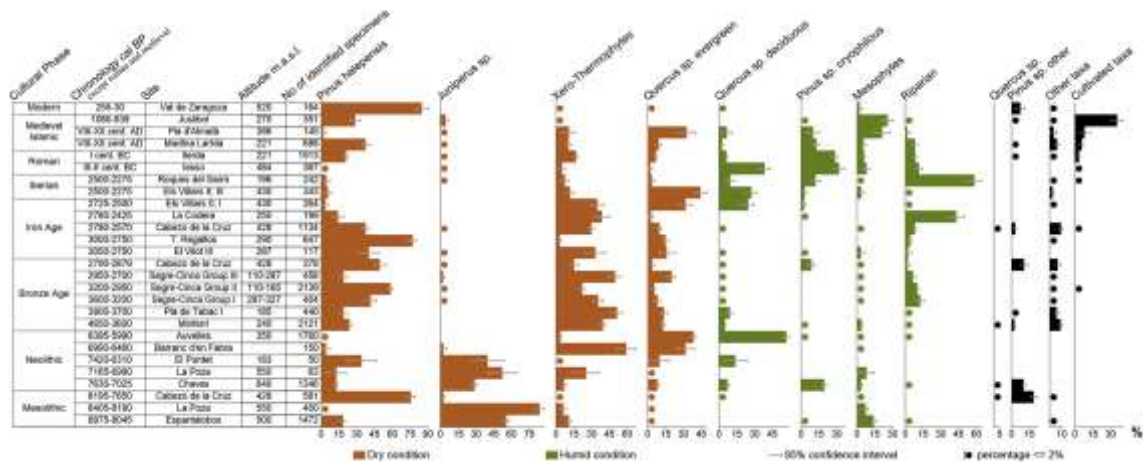


Fig. 4. Synthetic diagram showing main charcoal-based ecological groups throughout Mesolithic to Modern Ages. Raw data is presented in Tables 2, 4, 5, 6 and 7.

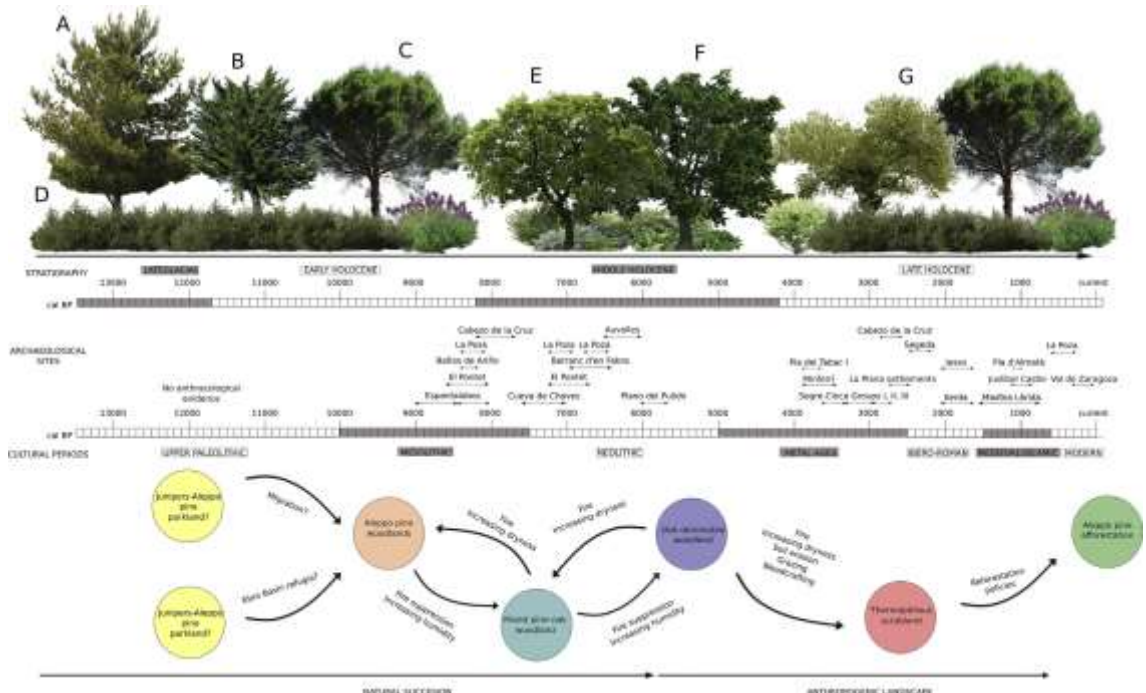


Fig. 5. Aleppo pine (*Pinus halepensis*) disturbance–succession model across different cultural/climate periods in the Ebro Basin. Mainwood charcoal-based landscape components have also been drawn: A) *Pinus nigra/sylvestris* type, B) *Juniperus* sp., C) *Pinus halepensis*, D) xero-thermophilous scrubland, E) *Quercus* sp. evergreen, F) *Quercus* sp. deciduous, G) *Olea europaea*. Segre-Cinca Group I includes Punta Farisa, El Vilot 0, Roques del Sarró; Segre-Cinca Group II comprises Genó, Masada de Ratón, Tossal de Solibernat and Vincamet I; Segre-Cinca Group III, Vilot I, II and Vincamet II. La Plana settlements refers to El Vilot III, Els Villars 0, I, La Codera and Tozal de Regalos Iron Age sites.