

Trend analysis of water quality series based on regression models with correlated errors

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Abstract

This work proposes a methodology for characterizing the time evolution of water quality time series taking into consideration the inherent problems that often appear in this type of data such as nonlinear trends, series having missing data, outliers, irregular measurement patterns, seasonal behavior, and serial correlation. The suggested approach, based on regression models with a Gaussian autoregressive moving average (ARMA) error, provides a framework where those problems can be dealt with simultaneously. Also the model takes into account the effect of influential factors, such as river flows, water temperature, and rainfall.

The proposed approach is general and can be applied to different types of water quality series. We applied the modeling framework to four monthly conductivity series recorded at the Ebro river basin (Spain). The results show that the model fits the data reasonably well, that time evolution of the conductivity series is non homogeneous over the year and, in some cases, non-monotonic. In addition, the results compared favorably over those obtained using simple linear regression, pre-whitening, and trend-free-pre-whitening techniques.

1 Introduction

2 The assessment of long term water quality trends is a subject of growing inter-
3 est. The 2000/60/EC directive of the European Parliament and of the Council
4 sets the goal of achieving a "good status" for all Europe's surface waters and
5 groundwater by 2015. According to the Water Framework Directive, Member
6 States must establish surveillance monitoring programmes to provide informa-
7 tion for the assessment of long-term changes in natural conditions and those
8 resulting from widespread anthropogenic activity. In this framework, the de-
9 tection and evaluation of the underlying trend due to anthropogenic activity
10 is a primary issue. The identification of periods and locations where increas-
11 ing pollution trends occur, would allow water management authorities to take
12 adequate measures.

13 Our interest is in providing a statistical method able to detect the presence
14 of a change in water quality, that can be attributable to anthropogenic be-
15 haviour. This is not an easy problem since those trends may be hidden by
16 the effect of external factors, such as river flow, seasonality, water tempera-
17 ture or precipitation. We search for an approach that estimates and extracts
18 the influence of these factors and, simultaneously, analyzes the presence of an
19 underlying temporal trend. The statistical analysis can not assure that this
20 trend is due to anthropogenic activities, but if the model properly eliminates
21 all the possible environmental factors, it can be assumed that the remaining
22 trend is related to human effect.

23 Alternative statistical methods such as non-parametric, time series and regres-
24 sion models have been used to assess trends in water quality (WQ in short)
25 series. In this work it is proposed a methodology which tries to solve the
26 disadvantages and limitations of those approaches. It enables us to take into
27 account the influence of non human factors and also to deal with common
28 problems in WQ series such as short records, missing data, outliers, irregular-
29 ity in the measurement pattern and, particularly, serial correlation. It is based
30 on the use of regression models with ARMA error and is successfully applied
31 to the analysis of four monthly conductivity series. This model could also be
32 used to simulate C behaviour under different environmental conditions. This
33 analysis was performed at the request of the Confederación Hidrográfica del
34 Ebro (CHE), the organization in charge of water management in the Ebro
35 river basin (Spain).

36 The paper is organized as follows. In section 2 the most common approaches
37 to the analysis of WQ series are reviewed, pointing out their advantages and
38 limitations with regard to the problem of interest. In sections 3 and 4 the
39 data and the model that are the basis of our approach are presented. Section
40 5 contains the core of the work, the description of how regression models
41 with ARMA error are used to carry out the trend analysis of conductivity
42 series and in section 6, the results for four conductivity series are discussed.
43 The comparison of these results with those obtained from other approaches is
44 shown in section 7 and the main conclusions are summarized in section 8.

46 In the WQ literature the most common approaches for detecting trends are
47 non-parametric. Mann-Kendall or other tests of association between two vari-
48 ables are used to detect the existence of monotonic temporal trends. Different
49 modifications have been made in order to deal with seasonality and serial cor-
50 relation, see for example Lettenmaier (1976), Hirsch et al. (1982) and Hirsch
51 and Slack (1984). Hirsch et al. (1991) examined some of the issues involved in
52 estimating WQ temporal trends. Hipel and McLeod (1994) Chap. 23, offer a
53 good review of the existing tests (seasonal Mann-Kendall, correlated seasonal
54 Mann-Kendall, Spearman partial rank correlation, etc.). A more recent review
55 on the identification of hydrological trends is provided by Khaliq et al. (2009).
56 This revision includes methods for incorporating the effect of serial correla-
57 tion, such as the pre-whitening approach proposed by Zhang et al. (2001),
58 the trend-free-pre-whitening technique by Yue et al. (2002), the variance cor-
59 rection approach and the block bootstrap. Other recent interesting references
60 on this topic are Hamed (2008), who developed a new version of the Mann-
61 Kendall test designed to account for the effect of scaling, and Bouza-Deaño et
62 al. (2008), who analyzed the WQ trends in the Ebro river.

63 The main disadvantage of non-parametric analysis is the difficulty of detect-
64 ing non monotonic trends. Moreover, most of the non-parametric tests do not
65 allow the evolution of WQ parameters to be linked to influential factors; Li-
66 biseller and Grimvall (2002) developed a partial Mann-Kendall test for trend
67 detection in the presence of covariates but it does not allow to model and
68 quantify the effect of those covariates.

69 Since WQ data are sequentially observed and often serially correlated, time
70 series techniques are used to model this kind of data. A wide review of time
71 series models can be found in Hipel and McLeod (1994) and numerous refer-
72 ences exist in WQ literature, such as Bhangu and Whitfield (1997), Worrall
73 and Burt (1999), Ahmad et al. (2001) and Lehmann and Rode (2001), for
74 example. However, time series models have some limitations, mainly that the
75 data must be observed at equally spaced time intervals and that these models
76 cannot easily deal with missing values, a common characteristic of WQ data.
77 Standard time series models can not easily take into account the influence of
78 external factors; Zetterqvist (1991) suggested the use of transfer functions to
79 include covariate series.

80 Regression modeling is another tool to assess time trends. It enables modeling
81 the effect of influential factors and it can easily deal with outliers, missing
82 observations and irregular measure patterns. A thorough review of this tech-
83 nique applied to WQ series can be found in Hipel and McLeod (1994) Chap.
84 24, and some more recent articles are Antonopoulos et al. (2001), Scarsbrook
85 et al. (2003) or Simeonova et al. (2003). More recently, Murdoch and Shan-
86 ley (2006) used segmented regression analysis to assess WQ trends while
87 Chang (2008) and Yang and Jin (2010) used spatial regression to incorpo-
88 rate the spatial correlation among the observations in the model estimation.
89 However, regression models are not often used in the WQ literature since their
90 assumptions (normality, constant variance, uncorrelation) are considered too
91 restrictive for usual WQ data. Regarding this point, it must be considered that
92 more general regression-based alternatives exist, such as the model suggested
93 in this work, whose assumptions can be more easily met by WQ data.

95 Electrical conductivity, also known as specific conductance and denoted by C
96 herein, is a frequently used WQ indicator. Our data set consists of monthly
97 conductivity series (microsiemens/cm) recorded at four rivers in the Ebro
98 basin (North-East of the Iberian Peninsula). The interest of WQ studies in
99 Mediterranean watersheds and in particular in the Ebro river basin is shown
100 in the special issue edited by Barceló and Sabater (2010) that includes pa-
101 pers on the particular problems of WQ and resource scarcity in regions with
102 a Mediterranean climate. The gauging stations were selected by CHE's man-
103 agers in order to consider rivers of different flow, water use, basin size, etc.,
104 see Figure 1. Data for the Ebro river were gauged at Tortosa, close to its
105 mouth, which represents the global behaviour of the basin; at Cinca river at
106 Fraga; at Segre river at Vilanova; and at Guatizalema river at Peralta. The
107 series were recorded from October 1980 to April 2003, except Peralta which
108 started in October 1981. The instantaneous river flow F (m^3/s) and the water
109 temperature WT ($^{\circ}C$) were also recorded.

110 Plots of the four C series are shown in Figure 2. The time points where the
111 three variables C , F and WT are available are joined by segments to show
112 the observations that could be used to fit a regression model with covariates
113 F and WT . A smoothed signal showing the long term evolution (a lowess
114 with a 35% window parameter) is also drawn. It can be seen that the lack of
115 complete observations, mainly due to the missing flow values, is an important
116 problem in Fraga and also in Vilanova and this problem makes the fit of a
117 good model more difficult. The lowess signals reveal an increasing long-term
118 trend in Tortosa while in Fraga and Vilanova after an increasing trend for the

119 first third of the record the series appear to fluctuate around a constant value;
120 in Peralta, a decreasing behaviour is observed from the late 90's. A thorough
121 preliminary analysis revealed that the series in Tortosa shows an unlikely
122 behaviour during the period 1988 to 1990, where all the observations are over
123 the smoothed mean level. To confirm the different behaviour of that period,
124 an indicator variable associated to the 3-year interval was introduced in the
125 model; since it was significant at a 0.01 level, these anomalous observations
126 were considered as missing values.

127 Some characteristics of the records are summarized in Table 1: the record
128 length, the number of missing observations in C , F and WT , the number of
129 complete observations in the three variables and the lag-1 correlation coef-
130 ficients of the monthly C series $\hat{\rho}_C(1)$. The serial correlations appear to be
131 significant and the highest values are observed for the Ebro (Tortosa) and
132 Cinca (Fraga) rivers, which have the largest flows.

133 The mean conductivity varies as 918 and 904 $\mu s/cm$ for Tortosa and Fraga,
134 and 550 and 503 $\mu s/cm$ for Vilanova and Peralta, respectively. The seasonal
135 means and standard deviations are shown in Figure 3, which suggest that the
136 series show some seasonality in the mean, although it is less evident in Peralta;
137 the standard deviations do not seem to show a seasonal behaviour.

138 The annual and seasonal mean values of the flow are summarized in Table 2.
139 This variable shows a clear seasonal behaviour with some common character-
140 istics in the four rivers: lowest mean values (in italic in the table) are always
141 observed in Summer, followed by the Autumn values; the highest mean val-
142 ues (in bold in the table) are observed in Winter except in Fraga, where it is
143 observed in Spring.

145 As aforementioned, ordinary regression model is a tool capable of dealing with
146 the usual problems in WQ data, except for the existence of serial correlation,
147 a common characteristic in monthly data. In all the conductivity series of this
148 study, residual correlation was detected after fitting an ordinary regression
149 model.

150 It is well known that with serially correlated data, the least squares estimator
151 is still unbiased but not BLUE (best linear unbiased estimator). Moreover, any
152 selection procedure applied during the modeling process will be doubtful since,
153 in the case of a positive correlation, the standard errors are underestimated
154 and the conclusions obtained from the t or F tests may be incorrect. Some
155 of the consequences of disregarding the error dependence in testing regression
156 coefficients were shown by Vinod (1976).

157 A solution to the correlation problem, which is often applied in Econometrics
158 and other fields, is some type of differentiation of the response variable, $y_t - y_{t-1}$
159 or $y_t - \rho(1)y_{t-1}$, where $\rho(1)$ is the lag-1 sample correlation coefficient. Other
160 possible solutions are the use of the lagged response as a covariate or the more
161 sophisticated ARMAX models, see Harvey (1990), that also include a linear
162 combination of the last values of covariate time series.

163 A drawback of all these approaches is that, in order to get the serial correlation
164 removed, the response variable is changed. The new response is a function of
165 Y_t and its past values and this makes it difficult to obtain the time evolution
166 of the original response. So, we opted for the use of a regression model with
167 ARMA error as an adequate and useful tool to obtain the underlying trend in

168 WQ series.

169 4.1 Regression models with ARMA error

170 The equation of a regression model with ARMA error is:

171

$$\mathbf{Y} = \mathbf{X}\beta + \mathbf{W} \tag{1}$$

172 where \mathbf{Y} is the vector of the response variable, \mathbf{X} the covariate matrix and \mathbf{W}

173 the error vector formed by values of a causal, zero-mean, ARMA(p,q) process,

174 satisfying

$$\Phi(B)W_t = \Theta(B)Z_t \tag{2}$$

175 where $Z_t \sim N(0, \sigma^2)$ is an uncorrelated series of normal random variables

176 with zero mean and constant variance, and $\Phi(B)$ and $\Theta(B)$ are polynomials

177 of order p and q in the backshift operator B , $BZ_t = Z_{t-1}$.

178 Estimation and inference in these models are well known (e.g. Brockwell and

179 Davis (2002) Chap. 6) and applications to various fields have been made.

180 Reinsel et al. (1981) and Reinsel et al. (1988) analyze stratospheric ozone

181 data for the detection of trends and Niu and Tiao (1995) generalize this

182 approach to a space-time regression model also for ozone data. In Medicine,

183 Greenhouse et al. (2006) develop a non-linear regression model with an ARMA

184 error for fitting biological rhythm data which they apply to series of human

185 core body temperature.

186 To our best knowledge this type of model has not been applied to WQ analysis.

187 In the next section the modeling process of WQ data using this approach is

188 presented, including a thorough description of its validation analysis; this is

189 an important and necessary step of any modeling process which must not be
190 underestimated.

191 5 Modeling process

192 5.1 The model

The regression model with ARMA error, succinctly described in (1), takes the following expression to represent the relationship between conductivity and the different covariates,

$$C_t = m_t(\beta^T) + f_1(WT_t; \beta^{WT}) + f_2(F_t; \beta^F) + s_t(\beta^s) + f_3(D_t; \beta^D) + f_4(t \times D_t; \beta^{DT}) + W_t$$

193 The vector of parameters β includes $\beta = (\beta^T, \beta^{WT}, \beta^F, \beta^s, \beta^D, \beta^{DT})$ where the
194 superscripts indicate the covariates. The elements of the model are:

- 195 - C_t is the conductivity at time t .
- 196 - $m_t(\beta^T)$ represents the time evolution, modeled by a polynomial, $\beta_1^T t + \beta_2^T t^2 +$
197 $\dots + \beta_k^T t^k$, whose order is determined during the modeling process.
- 198 - $f_1(WT_t; \beta^{WT})$ and $f_2(F_t; \beta^F)$ are, respectively, polynomial functions of wa-
199 ter temperature and river flow and, if necessary, in their lagged values, whose
200 orders are also fixed during the modeling process.
- 201 - $s_t(\beta^s)$ denotes a seasonal term modeled by a sum of Fourier harmonics. An-
202 other way of describing seasonal effects is by functions of indicator variables,
203 $f_3(D_t; \beta^D)$, linked to time periods such as seasons or months. Interaction
204 of these variables and time terms, $f_4(t \times D_t; \beta^{DT})$, are also considered in
205 order to enable the fit of different time evolutions in different periods of the
206 year. Due to their seasonal character, the fitted WT and F functions help

207 to model the seasonal behaviour of C .

208 - W_t are the error terms having an ARMA(p,q) structure, see equation (2).

209 *Invariance of the time definition.* The zero of the time variable t is located in
210 the middle of the record (January 1992) in order to avoid very high values,
211 and the time unit is one year; for example the time variable for Tortosa, a
212 series with 271 observations, is defined as $t = -135/12, -134/12, \dots, -2/12,$
213 $-1/12, 0, 1/12, 2/12, \dots, 135/12$. Since any other time origin and unit would
214 be equally adequate, a desirable property of the model would be its invariance
215 to possible origin and scale time changes. To achieve this property in models
216 including time interaction terms, a hierarchy principle must be applied: the
217 covariate in the interaction term must be included in the model even if it is not
218 statistically significant. For example, if D is not included, model (4) obtained
219 by a linear transformation of t in model (3), is not equivalent to model (3),

$$Y_t = \beta_0 + \beta_1 t \times D_t + W_t \quad (3)$$

$$\begin{aligned} Y_t &= \beta_0 + \beta_1 (a + bt) \times D_t + W_t \\ &= \beta_0 + \beta_1 a D_t + \beta_1 b t \times D_t + W_t \end{aligned} \quad (4)$$

220 However, if D is included, models (5) and (6) are equivalent,

$$Y_t = \beta_0 + \beta_1 t \times D_t + \beta_2 D_t + W_t \quad (5)$$

$$\begin{aligned} Y_t &= \beta_0 + \beta_1 (a + bt) \times D_t + \beta_2 D_t + W_t \\ &= \beta_0 + \beta_1 b t \times D_t + (\beta_1 a + \beta_2) D_t + W_t \end{aligned} \quad (6)$$

221 *Advantages of the use of a unified model with indicator variables.* The usual
222 approach in dealing with the seasonal character of WQ series is to fit separate
223 models for each season by splitting the sample. However, the suggested model
224 is able to simultaneously fit different structures for each season by using in-
225 dicator variables (binary variables that identify the observations of a season

226 or period of time) and their interactions with the covariates; see Weisberg
227 (2005), section 6.2, for more details. It results in a more complex model but
228 the set-up of a unified model has important advantages. First, it maintains
229 the sample size that, otherwise, would be divided; moreover, it is more flexible
230 since it allows us to consider not only the seasons but also other time periods
231 formed by one, or more, months. Nevertheless, the most important benefit is
232 that it provides a framework where the behaviour of the response in different
233 seasons or periods can be compared. These comparisons are carried out using
234 maximum likelihood ratio (MLR in short) tests that can objectively determine
235 if the time evolution of the response in two periods is the same. The merg-
236 ing of the similar periods, using adequate indicator variables, will result in a
237 reduction of the number of parameters and, consequently, in a simpler model.

238 5.2 *Estimating the linear predictor*

239 The parameter estimation is performed by restricted maximum likelihood,
240 REML in short, see Harville (1977). Ordinary maximum likelihood could
241 also be used but Cheang and Reinsel (2000) state that, for moderate sample
242 lengths, the REML estimator is, generally, less biased than the MLE and leads
243 to more accurate inference.

244 The covariate selection during the modeling process is mainly based on the
245 values of the t-ratios, $\hat{\beta}/\text{standard error}(\hat{\beta})$, associated to each covariate and, for
246 more complicated hypothesis, on the MLR test. This test allows checking any
247 hypothesis that involves the comparison of two models, one being a particular
248 case of the other (nested models); for example, the inclusion of the two terms
249 of a Fourier harmonic, the simultaneous effect of several terms, or the equality

250 of two coefficients of the model, $H_0 : \beta_i = \beta_j$. The significance level in all the
251 tests is 0.05.

252 The model selection is carried out in a systematic and iterative way, starting
253 at the simplest model and then comparing models in a stepwise approach,
254 until obtaining satisfactory fitting with the data. For each covariate, the hier-
255 archy principle is applied in the following way: a polynomial function with the
256 maximum order considered is first introduced; then, the greatest order term
257 is checked and removed from the model if it is non significant. This process
258 is repeated until a significant order term is found and then the simultaneous
259 effect of the current fitted polynomial function is checked using a MLR test.
260 Except when the sample size is too small, up to five order polynomials in time
261 and four order polynomials in WT and F are considered.

262 A common problem in WQ series is the presence of influential observations
263 (single cases or small group of cases that strongly influence the fit of the
264 model) and outliers (observations that fall unusually out of the pattern in
265 the relationship between the response and the covariates). In each step of
266 the covariate selection the existence of influential observations and outliers is
267 checked; the detected observations are removed from the sample to re-estimate
268 the model since our aim is to obtain a model which characterizes the vast
269 majority of the sample. The observations with a Cook distance higher than 1,
270 or much higher than the rest, are considered influential and the observations
271 with an absolute value of their standardized residuals much greater than 3 are
272 identified as outliers; see Weisberg (2005) Chap. 9 and Jobson (1991) Chap.
273 3 for details on this topic.

275 In each step of the covariate selection process, given the values p and q , the
276 coefficient vectors $\Phi = (\phi_1, \dots, \phi_p)$, $\Theta = (\theta_1, \dots, \theta_q)$ and β are estimated
277 simultaneously by REML.

278 In the first step, the selection of p and q is based on the serial correlation
279 structure of C , using the correlograms of the sample autocorrelation function
280 (ACF) and the partial autocorrelation function (PACF), together with indi-
281 vidual zero-correlation tests applied up to lag 24 (which corresponds to two
282 complete seasonal cycles). In the later steps, the same tools are used to test
283 if p and q values remain valid by analysing the residuals in that step. If there
284 is some evidence that they are no longer adequate, a MLR test is used to
285 compare the current model against the alternative suggested by the residual
286 dependence structure.

287 5.4 Validation analysis

288 The validation analysis aims to check if the assumptions of the model (the
289 relationship between the response and the covariates, the ARMA structure and
290 the normality of the error terms) are correct since, otherwise, the conclusions
291 obtained from the model could not be true.

292 *Residuals.* The validation analysis is based on the residuals e_t^* that result from
293 filtering the raw residuals $e_t = y_t - \hat{y}_t$ with the estimated ARMA process.
294 If the linear predictor and the ARMA structure are properly specified, the
295 filtered residuals must be zero-mean normal uncorrelated variables and the

296 usual regression diagnostics can be applied to them. When the ARMA error
297 reduces to an AR(p), the filtering is simple,

$$e_t^* = e_t - \hat{\phi}_1 e_{t-1} - \dots - \hat{\phi}_p e_{t-p}. \quad (7)$$

If the error structure is ARMA, the filtered residuals are

$$e_t^* = \hat{\Theta}(B)^{-1} \hat{\Phi}(B) e_t$$

298 and they should be calculated using for example a Kalman filter, see Brockwell
299 and Davis (2002).

300 *Checking serial correlation and normality.* To check if the filtered residuals
301 are uncorrelated, the correlograms of the ACF and PACF are plotted with an
302 approximate confidence band. Individual zero-correlation tests are also per-
303 formed up to lag 24 and, in addition, the Ljung-Box test (a portmanteau-type
304 test) is calculated for lags 6, 12, 18 and 24. Normality is checked using the
305 normal qqplot and the Shapiro-Wilk test.

306 *Checking the linear predictor and homoscedasticity.* The adequacy of fitted
307 predictors is checked using scatter plots of residuals versus the fitted values,
308 the covariates F , WT and time. In order to check if the seasonal behaviour is
309 adequately modeled, the same plots are drawn for each season and, in the case
310 of time, for each month. All these plots are also used to check homoscedasticity
311 (constant variance); this graphical tool is complemented with the Breusch-
312 Pagan test, see Cook and Weisberg (1993), which checks the hypothesis of
313 constant error variance against the alternative that variance changes with the
314 level of the fitted values or any covariate.

316 One of the advantages of regression models, including those with ARMA error,
317 is that their estimation and inference can be easily performed when some
318 observations are missing. As far as the validation analysis is concerned, the
319 tools for checking the normality and the predictor equation can deal with
320 missing observations but the analysis of serial correlation must be adapted.

321 *Testing serial correlation with missing observations.* The correlation function
322 $\rho(h)$ has to be estimated using only the complete pairs of observations. This
323 can significantly reduce the sample size, since each missing observation results
324 in the elimination of up to two pairs, and make the samples available for each
325 lag quite different. Because of this, the confidence intervals for $\rho(h)$ and the
326 Ljung-Box tests are calculated using the real sample size n_h instead of the
327 initial size n .

328 *Temporal trend analysis in records with long missing periods.* Another frequent
329 problem in WQ series is the existence of long periods without data. These
330 periods can make inadequate the use of a continuous polynomial function to
331 represent the time evolution. To overcome this problem, we opted for fitting
332 independent time polynomials in the isolated time periods. This is achieved
333 by fitting interactions of the time terms with indicator variables marking the
334 required periods. An example of this situation occurs in the Fraga series where
335 there is a ten year missing period, from 1988 to 1997, due to the lack of
336 river flow measurements. The use of two indicator variables linked to periods
337 1980-1988 and 1997-2003, respectively, allows fitting separate temporal trends
338 without reducing the sample size.

340 *6.1 Fitted models*

341 The most surprising result in the four fitted models was the complicated time
342 polynomials needed to represent the temporal evolution in Tortosa in Winter
343 (a 5-order polynomial function) and in Vilanova in Summer and Winter (a
344 4-order and 5-order polynomial functions, respectively), see Figure 4; in these
345 plots sample observations are plotted over the fitted curve to distinguish the
346 intervals where the estimation is based on fewer points. The evolutions suggest
347 that some important information, whose effect is mixed with time, is missing
348 in the linear predictors.

349 The non-monotonic behaviour of the temporal term fitted in Tortosa in Winter
350 follows quite approximately the cycles observed in the evolution of Winter
351 rainfall in the Eastern and Northeastern regions of the basin. The Winter
352 and Summer behaviour in Vilanova could also be explained by the effect of
353 the rainfall cycles in that region. So, two rainfall signals, *RainA* and *RainB*,
354 based on the anomalies of rainfall series observed in Tortosa-Tivissa and Lérida
355 regions, respectively, were calculated.

356 To check if regional rainfall is an influential factor, the residuals from the
357 models including only the significant covariates related to F and WT were
358 plotted versus the corresponding Winter or Summer rainfall signals. The plots
359 supported our hypothesis, see Figure 5 where Winter residuals for Vilanova
360 are plotted versus *RainB* together with a least square regression line. A new
361 stepwise selection procedure, starting from the previous final models without

362 the time terms and considering the rainfall covariates, was carried out for
363 Tortosa and Vilanova.

364 Some summary measures of the four final fitted models are outlined in Table 3:
365 the sample size, the number of outliers and influential observations eliminated
366 during the modeling process, the estimated white noise standard deviation $\hat{\sigma}$
367 and, as a goodness-of-fit measure, the square correlation coefficient between
368 the response and the fitted values, $Cor(C, \hat{C})^2$. The ARMA structure fitted
369 to the error terms and the corresponding $\hat{\Phi}$ coefficients are shown in the last
370 two rows. The covariates included in each linear predictor together with their
371 coefficients and the p-values of their t-ratios are shown in Tables 4 (Fraga and
372 Peralta) and 5 (Tortosa and Vilanova). It must be remembered that some non
373 significant covariates according to t-test are kept in the model if they make
374 part of a significant simultaneous effect or following the hierarchy principle,
375 see section 5.2. The season indicator variables are denoted by Sp (Spring), Su
376 (Summer), Au (Autumn) and Wi (Winter) and the variables associated with
377 the months by the first three letters of their name, e.g. February is denoted by
378 Feb . The fitted temporal trends for the four final models are shown in Figure
379 6.

380 The new model for Vilanova is quite simpler. The introduction of $RainB$ in
381 Winter and Summer reduces the previous 5 and 4-order time polynomials to
382 simple linear terms. Initially, two different slopes were fitted for Summer and
383 Winter periods, 6.62 and 6.86 ($\mu s/cm$) $year^{-1}$ respectively, but finally both
384 seasons were jointly modeled since a MLR test leads to conclude that there is
385 not a significant difference between the slopes at a 0.05 level.

386 The model for Tortosa after adding the rainfall term is also simpler. The

387 introduction of the *RainA* signal in Winter reduces the previous five order
388 time polynomial to a linear term with slope $11.24 (\mu s/cm)year^{-1}$. The tem-
389 poral slope for the Spring-Summer-November period in this model is 13.65
390 $(\mu s/cm)year^{-1}$ and since a MLR test does not reject the slope equality at the
391 0.05 level, the trends of both periods are jointly fitted. Hence, there is a linear
392 trend in C except for the months of September and October, where no tem-
393 poral evolution is detected. It is noteworthy that although the $Wi \times RainA$
394 and Wi terms in the Tortosa model, with p-values equal to 0.08 and 0.10
395 respectively, are not significant according to t-tests, the simultaneous effect
396 of these two covariates is highly significant, with the p-value of the MLR test
397 lower than 10^{-6} .

398 The period 1997-2003 in Fraga cannot be analyzed seasonally due to the scarce
399 number of observations.

400 *6.1.1 Comparing the fitted models*

401 The river flow influences C in the four fitted models, although in different ways:
402 linearly in Peralta, quadratically in Vilanova and as a third order polynomial
403 in Fraga. The lagged flow is influential only in Tortosa where up to the second
404 lag is significant. The greater complexity of this model is due to the fact that
405 large rivers have much more persistence than do small streams and probably
406 favored by the bigger sample size. The effect of WT is linear and positive in
407 Fraga and Vilanova and also linear, but negative, in Tortosa. In Peralta, the
408 temperature effect is quadratic.

409 As to seasonality, no harmonic term is included in any of the models since
410 the seasonal effect is better described by F , WT and indicator variables. The

411 fitted models show that a lower conductivity, with respect to the mean level of
412 the series, is observed in Summer in Peralta, Vilanova and in Fraga during the
413 period 1980-1987, with coefficients -52.7, -32.8 and -80.2 $\mu s/cm$, respectively.

414 On the other hand, a higher conductivity level is observed in Vilanova in
415 Winter and in Tortosa in the period AuDec (Autumn and December), with
416 coefficients 106.0 and 216.1 $\mu s/cm$ respectively.

417 The fitted models reveal that the temporal evolution of C is not homogeneous
418 throughout the year. In Vilanova there are linear increasing trends, with a
419 common slope, in Summer and Winter, but C does not show any significant
420 change in Spring nor in Autumn. In Tortosa there is also an increasing linear
421 trend during all the year, except for September and October. In Peralta, an
422 increasing linear trend is found in December and a similar evolution is observed
423 between 1986 and 1997 in Spring and Autumn; however, the global evolution
424 in these seasons is non monotonic and decreases from 1997. A decreasing linear
425 trend from 1997 onwards is also detected in Fraga.

426 The observed increase of conductivity is not surprising since the use of water
427 for agricultural purposes has grown in the Ebro basin during recent decades.
428 The irrigated surface increased from 650 000 Ha in 1980 to 700 000 Ha in 2002.
429 Concerning the simultaneous decreasing trends observed from 1997 onwards
430 in the Guatizalema (Peralta) and Cinca (Fraga) rivers, a sound and common
431 explanation could be the following. The Guatizalema basin, upstream of Per-
432 alta, is irrigated with water coming from the Cinca river through the Cinca
433 channel, which starts downstream the system formed by El Grado and Medi-
434 ano reservoirs. This system suffered an important and monotonic decrease of
435 its water supply during the period 1995-2005. Apparently, if there is less water
436 available, the farmers grow crops which demand less water and fertilizer and,

437 given the lesser water return to the river, these crops are less pollutant. This
438 explanation, provided by CHE experts, cannot be confirmed since agrarian
439 surveys about crop history in these areas are not available.

440 With regard to the serial correlation, the fitted structure in the four models is
441 an autoregressive process, which reduces to a simple AR(1) except for Peralta
442 model.

443 *6.2 Validation analysis*

444 The four models have been fully validated with respect to serial correlation
445 structure, homoscedasticity and the linear predictor specification (a possible
446 misrepresentation of the covariates in the model formula or the absence of
447 relevant terms). The normality of the residuals is also checked since it is a
448 necessary assumption to make inference.

449 *Serial correlation and normality.* Table 6 shows the p-value of the lag-1 au-
450 tocorrelation test for e_t^* and the number of lags, between 1 and 24, with a
451 significant coefficient at the usual 0.05 level (the significant lags are shown
452 in brackets). The p-values of the Ljung-Box test for lags 6, 12, 18 and 24,
453 are also given and the ACF and PACF correlograms for Tortosa are shown
454 in Figure 7 as an example. All the results provide a strong evidence on the
455 uncorrelation of the filtered residuals. According to the normal qqplots and
456 the Shapiro-Wilk test, see the p-values in Table 6, the normality of the filtered
457 residuals can not be rejected in any model.

458 *Linear predictor and homoscedasticity.* The scatter plots of filtered residuals
459 versus fitted values, water temperature, river flow and time are shown for the

460 Tortosa model in Figure 8; the linear predictor seems to be well specified since
461 no pattern is detected in these plots. Concerning heteroscedasticity, although
462 the variability seems to increase with the fitted values, see plot bottom right,
463 the Breusch-Pagan test shows that this increase is not significant, see the p-
464 value in Table 6. The seasonal structure also seems to be adequately modeled,
465 see Figure 9. The same analyzes for the remaining models also turn out to be
466 satisfactory.

467 6.3 Software

468 The modeling process is performed using the freeware R package, which can be
469 downloaded from The Comprehensive R Archive network (CRAN) at [http://cran.r-](http://cran.r-project.org/)
470 [project.org/](http://cran.r-project.org/). The model estimation is carried out with the function *gls* from
471 the library *nlme*. Some own functions and the package *car* are used in the
472 validation analysis.

473 7 Comparison with other approaches

474 In order to compare the performance of the ARMA regression approach, the
475 four *C* series were analyzed using other common approaches: a simple linear
476 (SL) regression versus time, the pre-whitening (PW) approach proposed by
477 Zhang et al. (2001) and the trend-free-pre-whitening (TFPW) technique by
478 Yue et al. (2002). In the PW technique, if the lag-1 serial correlation $\hat{\rho}(1)$ is
479 significant, the trend test is applied to the pre-whitened series $(y_t - \hat{\rho}(1)y_{t-1})$.
480 Since this procedure affects the magnitude of the true slope in the original
481 series, the TFPW tries to overcome this problem by calculating $\hat{\rho}(1)$ from the

482 series detrended using the Sen slope estimator and, if the serial correlation
483 coefficient is significant, the trend test is applied to the detrended pre-whitened
484 series recombined with the Sen slope.

485 In Vilanova and Peralta, where the temporal evolution found by the ARMA
486 regression approach is not the same in the different seasons, the three alterna-
487 tive methods are applied one season at a time. In Tortosa, since the ARMA
488 regression finds a common time behaviour, except in September and October,
489 the three methods are applied to the complete series. In Fraga two separate
490 analysis for periods P1 and P2 are performed. The slopes estimated using the
491 four approaches and their corresponding p-values are shown in Table 7. In all
492 the cases, the null hypothesis is that there is no linear temporal trend, i.e.
493 that the slope is equal to zero.

494 In Tortosa, the four approaches detect a significant linear temporal trend
495 whose slope is underestimated if factors such as F and WT are not properly
496 taken into account. The slope decreases from 13 with the ARMA regression to
497 $8.1(\mu s/cm)year^{-1}$ with the TFPW method. In addition, the ARMA regression
498 detects that this trend is not homogeneous over the whole year and identifies
499 two months, September and October, with no trend.

500 In Vilanova, the four methods provide the same result for Spring and Au-
501 tumn where no temporal trend is detected. In Summer and Winter, only the
502 ARMA regression is able to detect an underlying positive temporal trend,
503 after eliminating the wavy rainfall effect.

504 In Peralta, the ARMA regression detects a temporal trend in Autumn and
505 Spring, described by a cubic polynomial, which decreases from about 1997
506 onwards. The other approaches, designed to identify monotonic trends, do

507 not detect this behaviour. In Winter and Summer no trend is found by the
508 methods that take into account the existing serial correlation. The possible
509 mistakes resulting from not accounting for the serial correlation are evident
510 in Winter, where the simple regression finds a statistically significant positive
511 temporal trend.

512 In the P1 period of Fraga, the PW and the TFPW methods detect a pos-
513 itive trend, both with a p-value equal to 0.07, while the ARMA regression
514 does not. These results are not comparable since the series used in the ARMA
515 regression has 28 observations less, from a sample of 106, due to the lack of
516 flow data. When the estimation and tests are applied to the same series, the
517 PW and the TFPW methods do not find any significant trend, see values in
518 brackets in Table 7. For the P2 period, 1998-2003, only the ARMA regres-
519 sion is able to detect a significant negative linear temporal trend in the C
520 behaviour. The explanation of this fact is that the river flow shows a decreas-
521 ing evolution during this period. The effect of a decreasing flow compensates
522 the existing decreasing trend of conductivity and makes the PW and TFWP
523 methods underestimate the real negative C trend with values -55.4 and -21.8
524 $(\mu s/cm)year^{-1}$ respectively, that are not significant. On the other hand, the
525 temporal slope estimated by the ARMA regression, which takes into account
526 the F effect, is $-90.4 (\mu s/cm)year^{-1}$ and significant.

527 The fit of temporal trends using the PW and the TFPW non-parametric pro-
528 cedures was carried out using the *zyp* library in R, developed by D. Bronaugh
529 and A. Werner.

531 It can be concluded that the use of a regression model with a Gaussian ARMA
532 error is a powerful tool to characterize, quantify and make inference on the
533 temporal evolution of water quality parameters, under the difficult conditions
534 often found in these series. Although the modeling process can be more time-
535 consuming than other simpler approaches, it has important advantages as
536 summarized below.

- 537 • It deals with the limitations of other common approaches to WQ trend
538 analysis. More precisely,
 - 539 - It can be used in series with a serial correlation structure.
 - 540 - Because the proposed approach can characterize complex temporal evo-
541 lutions such as non-monotonic trends, it offers an advantage over non-
542 parametric trend analysis
 - 543 - It can be easily applied to series with an irregular recording pattern and
544 missing observations. The Fraga series is an example of how it can be
545 applied to data having long missing periods.
 - 546 - It can take into account any available influential factor.
- 547 • ARMA regression models are a flexible tool to analyze any temporal evolu-
548 tion. In particular,
 - 549 - The use of maximum likelihood ratio tests in combination with adequate
550 indicator variables allows modeling different trends in different seasons,
551 months or groups of months. It also allows describing these different evo-
552 lutions in the simplest way by testing whether different periods have iden-
553 tical parameters.
 - 554 - The seasonal behaviour can be described by different and flexible tools:

555 harmonic terms, covariates with a seasonal character or indicator variables
556 linked to the most convenient periods.

557 - Using appropriate tests it is possible to determine which factors are influ-
558 ential and in which way they influence the response: linearly, through a
559 polynomial function, through past values, etc.

- 560 • This approach uses the available information more efficiently, since the
561 whole sample is used to estimate the model, whereas in other methods
562 the sample must be split to undertake seasonal analysis or to distinguish
563 different periods.
- 564 • The model validation process helps to improve the model in an iterative
565 way since the residual analysis not only detects its failures but also suggests
566 how to solve them. In our case, it helped detecting the influence of rainfall,
567 a factor not considered initially.
- 568 • If influential natural factors, such as river flow, water temperature or rain-
569 fall, are not properly considered in the modeling process, as the ARMA
570 regression does, the temporal trends due to anthropogenic activities can
571 be hidden or distorted. The series of Vilanova in Winter and Summer and
572 Tortosa are clear examples of this possibility.

573 As regards the results obtained in the four gauging stations in the Ebro basin,
574 it is found that the temporal evolution of C series is not homogeneous within
575 the year and sometimes even within seasons. The behaviour of C is not spa-
576 tially homogeneous, although most of the series are stable or show an in-
577 creasing trend most of the time; only Peralta and Fraga show a decreasing
578 conductivity trend from about 1997. These differences are not surprising since
579 the analyzed rivers present quite different characteristics regarding their flow,
580 basin area and use of the water. From the six significant trends detected, five

581 are linear and one is non-monotonic.

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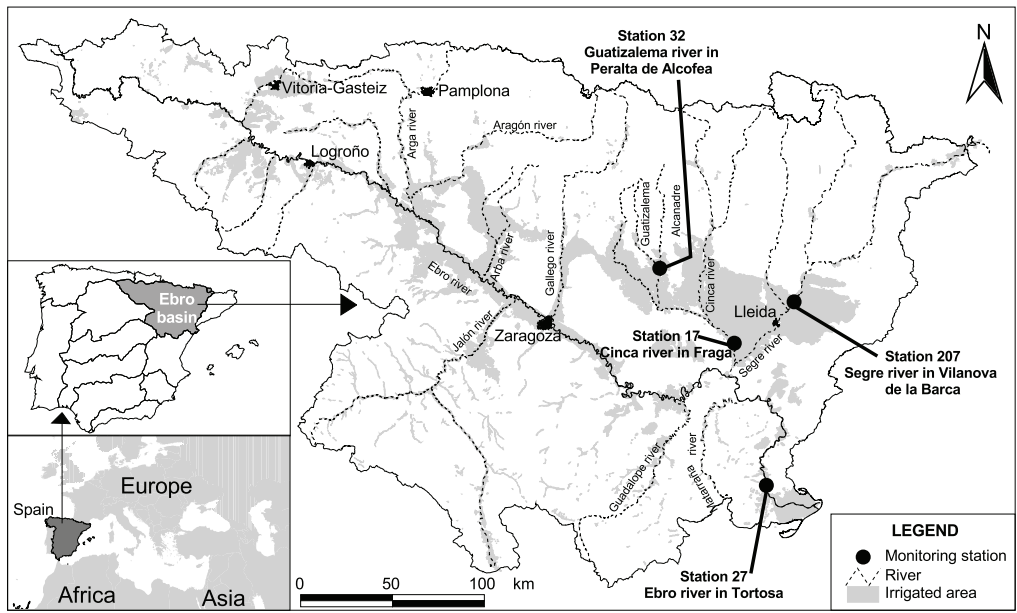


Fig. 1. Location of the Ebro river basin and the gauging stations.

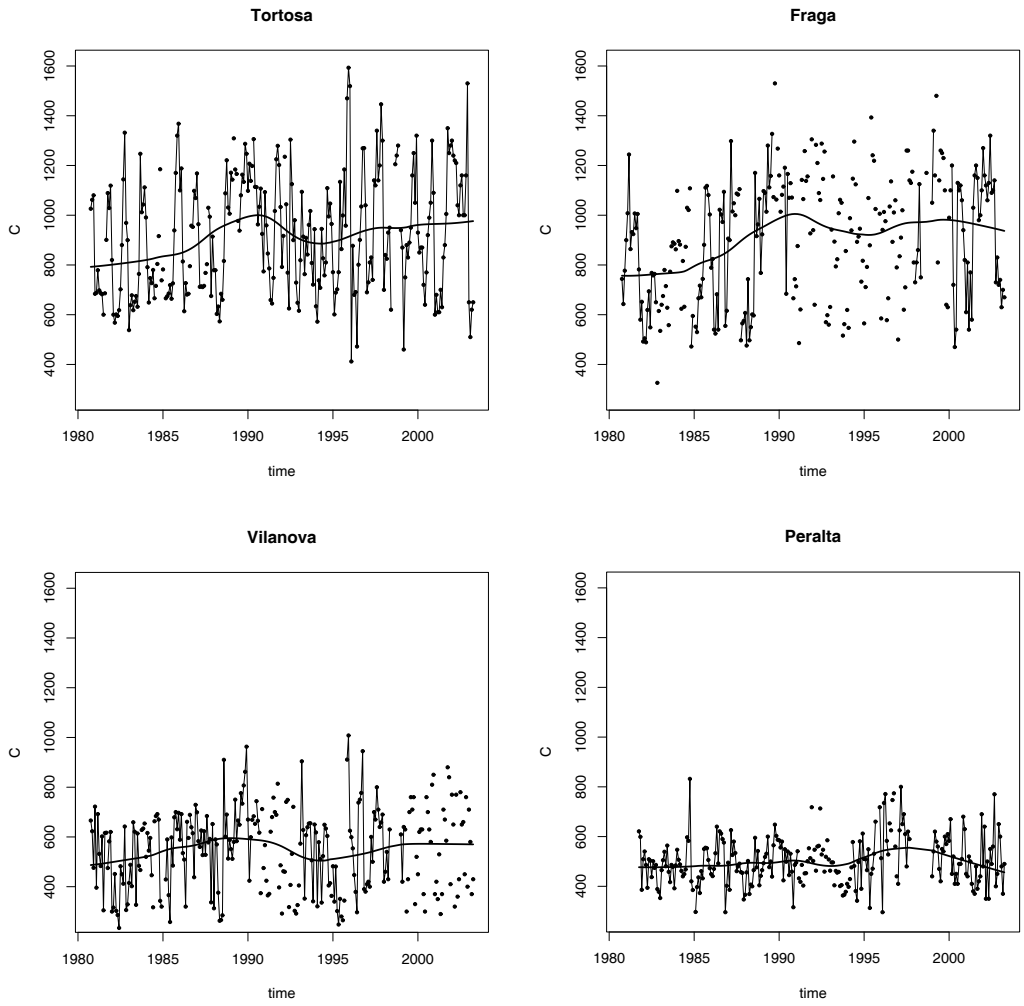


Fig. 2. Monthly series of conductivity (points) and long term evolution signals (lowess smoother). The segments joining the points indicate the periods where the three variables C , F and WT are available.

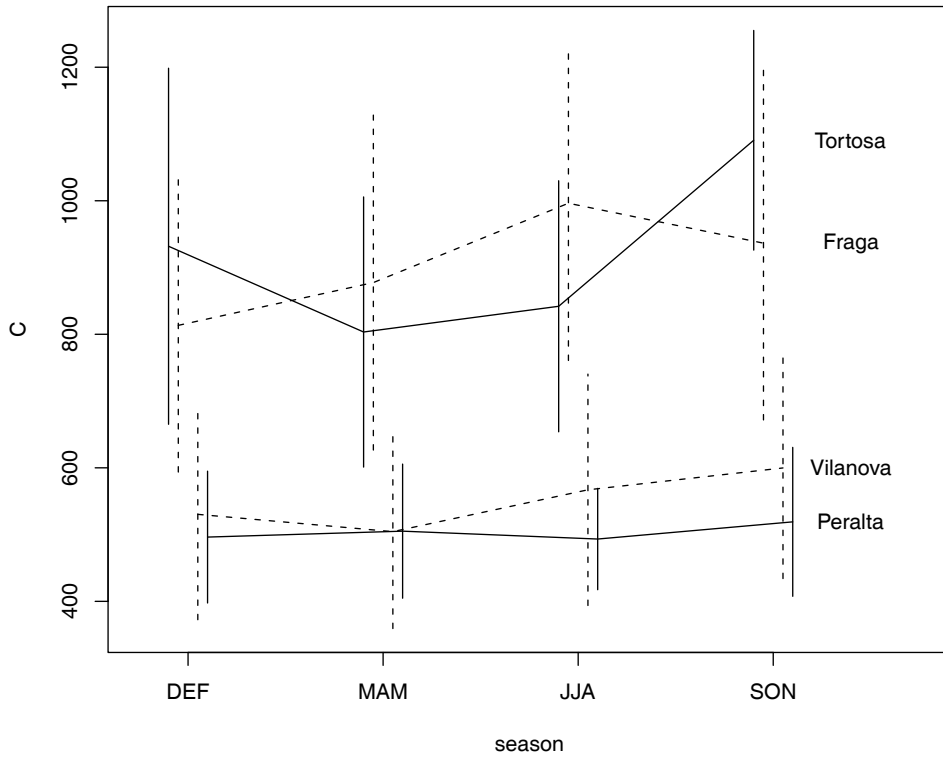


Fig. 3. Seasonal means \pm standard deviation for the C series.

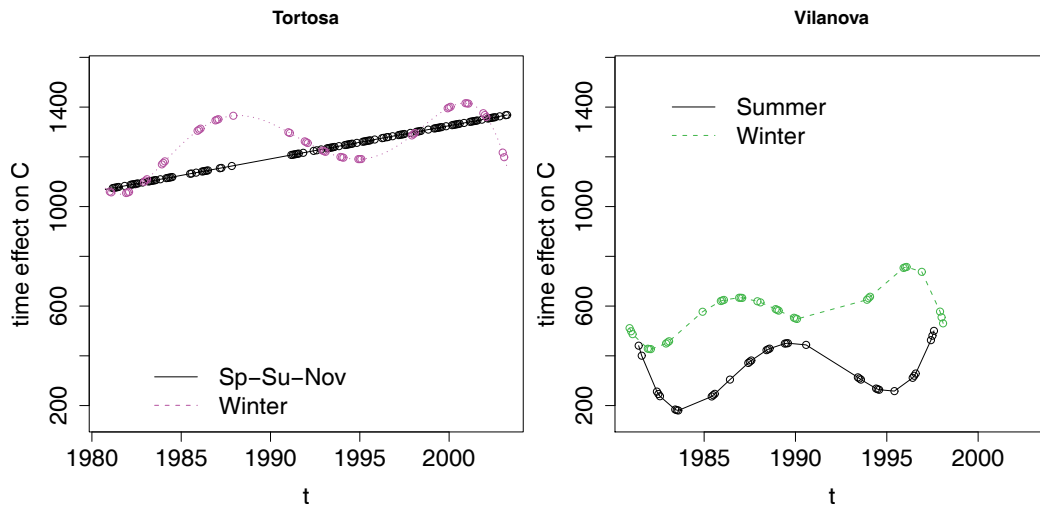


Fig. 4. Time evolutions fitted in the ARMA regression models (without rainfall signals), Tortosa and Vilanova.

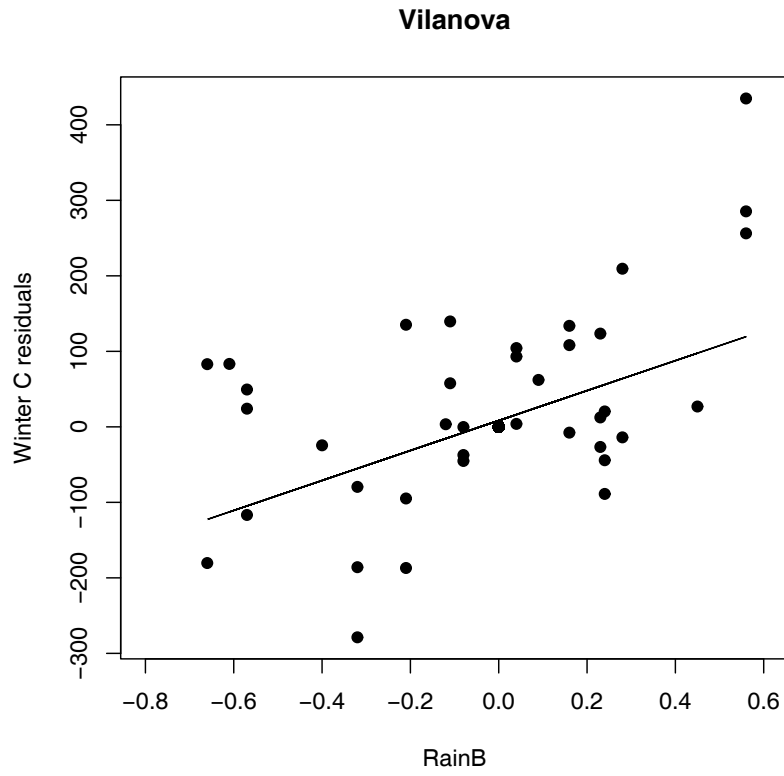


Fig. 5. Scatter plot of Winter conductivity residuals for Vilanova, from a model including only F and WT effects, versus the regional rainfall signal.

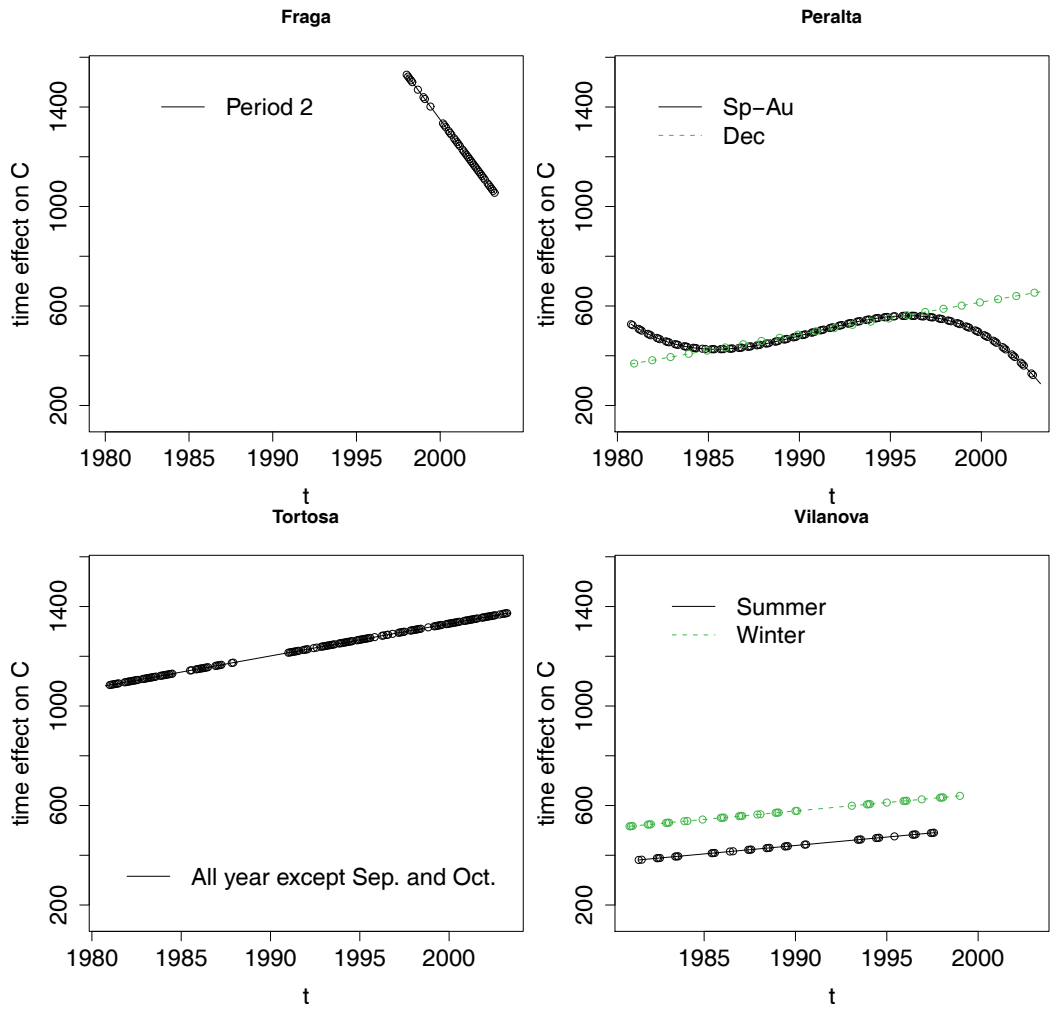


Fig. 6. Time evolutions fitted in the final ARMA regression models for the four series.

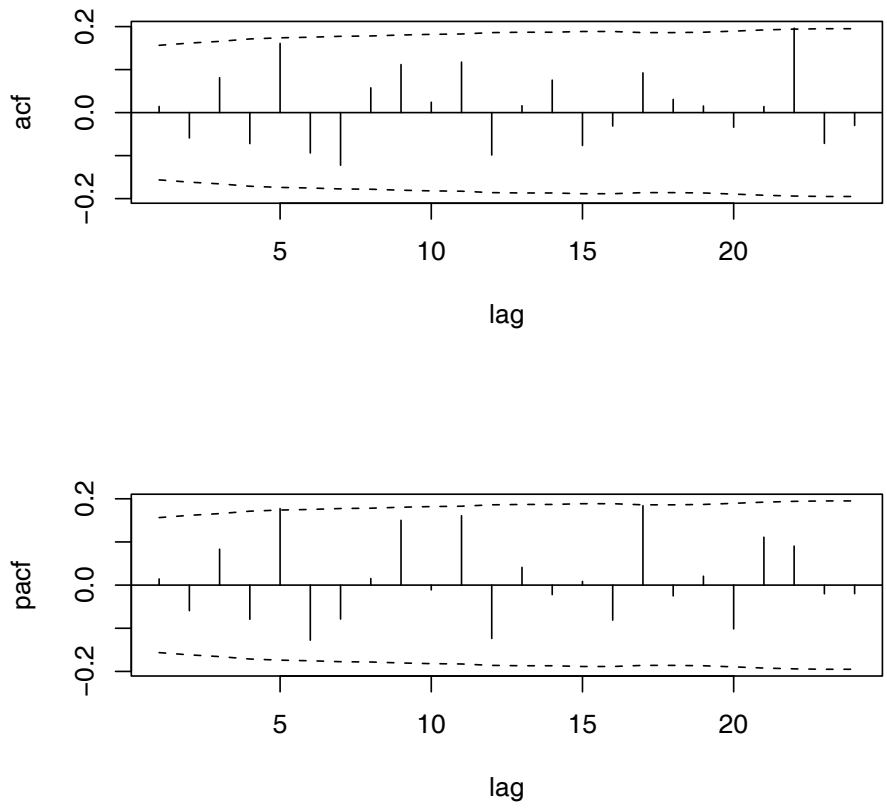


Fig. 7. Correlograms of the ACF and PACF of filtered residuals from the Tortosa model.

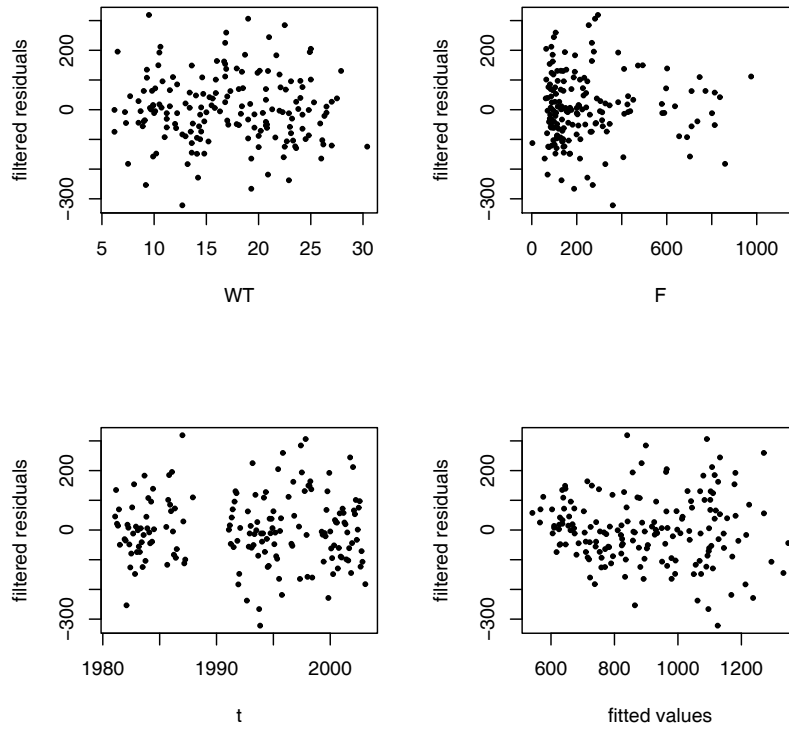


Fig. 8. Scatter plots of the filtered residuals versus fitted values and the covariates, water temperature, river flow and time for the Tortosa model.

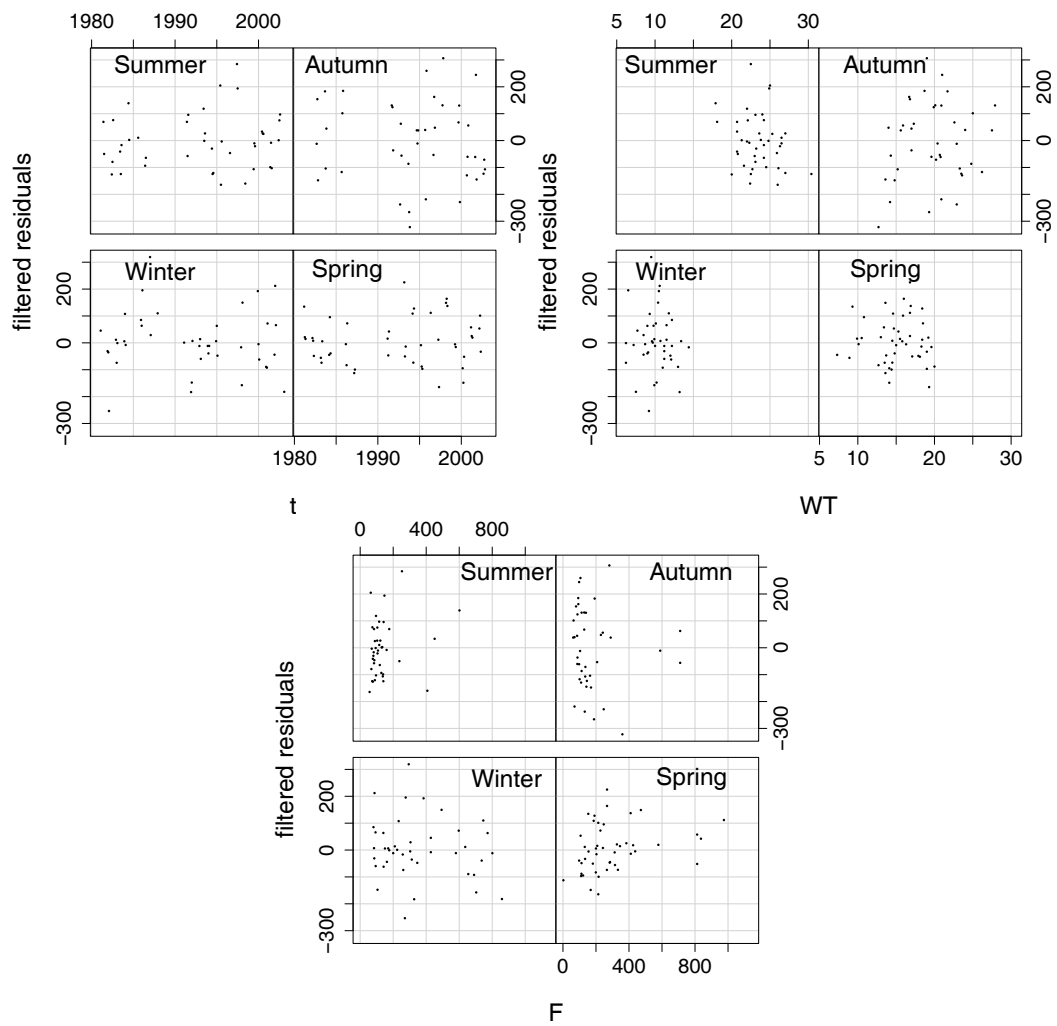


Fig. 9. Seasonal scatter plots of filtered residuals versus time, water temperature and river flow for Tortosa.

677 Table captions

678 Table 1. Some characteristics of the records of the four analyzed series.

679 Table 2. Annual and seasonal flow mean values; lowest seasonal values for each
680 location are in italic and highest in bold.

681 Table 3. Summary measures of the fitted ARMA regression models.

682 Table 4. Coefficients and p-values of fitted linear predictors in Fraga and Per-
683 alta models.

684 Table 5. Coefficients and p-values of fitted linear predictors in Tortosa and
685 Vilanova models.

686 Table 6. Summary of some validation measures based on e_t^* for the four final
687 models.

688 Table 7. Estimated time slopes ($\mu s/cm$) $year^{-1}$ and p-values corresponding to
689 null slope tests using four different approaches: the SL regression, the PW
690 method, the TFPW method and the ARMA regression.

	Tortosa	Fraga	Vilanova	Peralta
Record length	271	271	271	259
# missing obs. in C	40	9	10	18
# missing obs. in F	18	143	100	59
# missing obs. in WT	4	10	10	18
# complete obs.	216	127	171	200
$\hat{\rho}_C(1)$	0.64	0.55	0.32	0.35

Table 1

	Tortosa	Fraga	Vilanova	Peralta
Annual mean F (m^3/s)	260.5	60.5	21.4	1.2
Winter mean F	388.7	61.1	30.0	2.8
Spring mean F	308.9	68.9	22.0	0.8
Summer mean F	<i>149.2</i>	<i>54.1</i>	<i>15.8</i>	<i>0.5</i>
Autumn mean F	184.2	56.6	16.6	0.6

Table 2

Series	Fraga	Peralta	Tortosa	Vilanova
Sample size	116	190	185	159
# influential obs.	5	8	10	7
# outliers	6	2	2	5
$\hat{\sigma}$	126.4	73.7	126.1	115.8
$Cor(C, \hat{C})^2$	0.76	0.42	0.71	0.48
error ARMA struct.	AR(1)	AR(2)	AR(1)	AR(1)
$\hat{\Phi}$	0.49	0.19, 0.27	0.29	0.20

Table 3

Fraga	$\hat{\beta}$	p-value
const.	1302.6	0.00
F	-18.1	0.00
F ²	0.1	0.00
F ³	0.0004	0.00
WT	10.3	0.00
P2× <i>t</i>	-90.4	0.00
P2	769.8	0.00
Su× P1	-80.2	0.04

Peralta	$\hat{\beta}$	p-value
const.	481.4	0.00
F	-37.7	0.00
WT	9.2	0.02
WT ²	-0.2	0.07
AuSp× <i>t</i>	17.9	0.00
AuSp× <i>t</i> ²	-0.9	0.00
AuSp× <i>t</i> ³	-0.2	0.00
AuSp	36.0	0.10
Dec× <i>t</i>	12.9	0.00
Dec	30.7	0.10
Su	-52.7	0.04

Table 4

Tortosa	$\hat{\beta}$	p-value
const.	1191.8	0.00
F	-0.3	0.00
lag F	-0.4	0.00
lag2 F	-0.7	0.00
lag2 F ²	0.0005	0.00
WT	-6.9	0.01
SpSuWiNov $\times t$	13.0	0.00
SpSuWiNov	35.6	0.30
Wi \times RainA	140.7	0.08
Wi	53.4	0.10
AuDec	216.1	0.00

Vilanova	$\hat{\beta}$	p-value
const.	485.6	0.00
F	-5.6	0.00
F ²	0.03	0.00
WT	10.3	0.00
Su \times RainB	116.2	0.08
Su	-32.8	0.34
Wi \times RainB	173.7	0.00
WiSu $\times t$	6.8	0.01
Wi	106.0	0.00
May	-114.6	0.00

Table 5

Series	Tortosa	Fraga	Vilanova	Peralta
p-value of the correlation test	0.85	0.53	0.23	0.45
# signif. ACF lags (signif. lags)	1 (22)	1 (13)	0	1 (9)
Ljung-Box p-value: 6 lags; 12 lags	0.21; 0.19	0.13; 0.21	0.48; 0.25	0.74; 0.46
18 lags; 24 lags	0.39; 0.35	0.12; 0.08	0.25; 0.43	0.65; 0.66
Shapiro-Wilk p-value	0.50	0.10	0.19	0.96
Breusch-Pagan p-value	0.46	0.65	0.21	0.68

Table 6

Series	SL regression		PW method		TFPW method		ARMA regression	
	Slope	p-value	Slope	p-value	Slope	p-value	Slope	p-value
Tortosa	7.2	0.00	9.9	0.01	8.1	0.01	13.0	0.00
Vilanova, Spring	1.5	0.67	0.9	0.86	1.6	0.84	-	-
Vilanova, Summer	7.1	0.16	6.3	0.31	7.7	0.31	6.8	0.01
Vilanova, Autumn	2.6	0.57	-1.4	0.77	2.6	0.79	-	-
Vilanova, Winter	-0.5	0.91	-6.1	0.35	-2.1	0.36	6.8	0.01
Peralta, Spring	1.1	0.57	1.7	0.68	0.7	0.69	cubic poly.	
Peralta, Summer	-0.4	0.79	-0.4	0.92	-0.1	0.92	-	-
Peralta, Autumn	1.3	0.59	1.59	0.8	0.5	0.80	cubic poly.	
Peralta, Winter	6.8	0.00	0.67	0.8	6.8	0.53	-	-
Fraga, P1	22.7	0.01	26.2	0.07	24.9	0.07	-	-
(obs. with F value)	(17.0)	(0.07)	(24.4)	(0.24)	(18.7)	(0.22)	-	-
Fraga, P2	-19.4	0.37	-55.38	0.3	-21.8	0.26	-90.4	0.00

Table 7